

# Intensive smooth cordgrass removal strengthens tidal and temperature impacts on methane emission<sup>☆</sup>

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## ABSTRACT

The world's largest ecosystem restoration via intensive removals of invasive smooth cordgrass (*Spartina alterniflora*) is being implemented in coastal China, potentially exerting a large impact on soil biogeochemical cycles of greenhouse gases including methane (CH<sub>4</sub>). However, the degree to which CH<sub>4</sub> emission and its environmental controls change with such anthropogenic disturbances has been rarely assessed with direct empirical evidence. To quantify these disturbance effects, we utilized the eddy covariance (EC) approach to continuously measure net CH<sub>4</sub> exchange from Jul. 2022 to Oct. 2023, covering both pre- and post-removal periods, in a disturbed coastal wetland of Southeast China experiencing an intensive cordgrass removal in late Oct. 2022. Our analyses, based on this unique EC dataset of high-frequency (30-min) time series CH<sub>4</sub> fluxes, revealed that (a) the removal caused a pulse of CH<sub>4</sub> emission peaking one month later up to 0.76 g CH<sub>4</sub> m<sup>-2</sup> d<sup>-1</sup>, with the mean post-removal emission over ten times that of the pre-removal level (0.03 g CH<sub>4</sub> m<sup>-2</sup> d<sup>-1</sup>); (b) the removal intensified the controls of tidal inundation and pumping on CH<sub>4</sub> fluxes, showing much stronger pumping effects within two months following the disturbances; (c) the removal also enlarged the temperature sensitivity of CH<sub>4</sub> emission, leading to larger daytime emission especially at afternoon hours; (d) the combination of enhanced tidal impacts and temperature dependence thus promoted the diel variability of CH<sub>4</sub> fluxes during the post-removal period. These results suggest that coastal restoration via intensive cordgrass removals boosts both the magnitude and the diel variability of CH<sub>4</sub> emission, highlighting the necessity of better understanding the climate impact of restoration activities. Future longer flux data with extended years are needed to further assess potential regime shift in soil CH<sub>4</sub> biogeochemistry and long-term evolution of such unintended environmental costs of the restoration.

## 1. Introduction

As a cost-effective natural solution for mitigating climate change, coastal blue carbon ecosystems (BCEs) have attracted significant attentions due to their efficient carbon sequestration capacity (Macreadie et al., 2021; Wang et al., 2023a). BCEs can reduce approximately 304 Tg carbon dioxide (CO<sub>2</sub>) emission annually (Macreadie et al., 2021). However, as hotspots of biogeochemical cycling of greenhouse gases (GHGs), they also emit 0.47–1.41 Tg methane (CH<sub>4</sub>) annually (Rosentreter et al., 2023), significantly offsetting their carbon

sequestration benefits (Rosentreter et al., 2018). Unfortunately, the contribution of CH<sub>4</sub> emission is often overlooked in current blue carbon assessments. More critically, factors such as biological invasions, anthropogenic land-use change, and climate change are accelerating the degradation of BCEs (Macreadie et al., 2021). As one of the world's most threatened habitats, BCEs are being lost at an annual rate of 0.13–1.5%, resulting in annual emission of 0.15–1.02 Pg CO<sub>2</sub> and severely compromising their carbon sequestration potential (Fu et al., 2024). In response to the United Nations Decade on Ecosystem Restoration, urgent and effective measures must be taken to enhance the protection and

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restoration of BCEs, maximizing their carbon sequestration capacity and ecosystem service benefits.

China occupies one of the most severely degraded coastal zones in the world, and has implemented systematic ecological restoration measures (e.g., invasive species removal and habitat reconstruction) to strengthen the protection of BCEs (Liu et al., 2016). Originally native to the Atlantic coast of the United States, smooth cordgrass (*Spartina alterniflora*) has spread rapidly since its introduction to China in 1979 and occupies over 60% of the country's total saltmarsh area along the coastline (Nie et al., 2023). Recently, the Chinese government launched a large-scale removal campaign in 2022, aiming to eliminate 90% of the invasive cordgrass by 2025 (Stokstad, 2023; Xie and Han, 2023). Recent studies have primarily used scenario analyses to assess the long-term ecological benefits of this campaign (Qi et al., 2024; Yang et al., 2024), but insufficient attention has been paid to the short-term ecological effects of such intensive anthropogenic disturbances. In fact, wetland ecosystems are highly sensitive to anthropogenic disturbances (He et al., 2025). Human-induced land use and land cover changes can significantly alter wetland landscape patterns, thereby affecting their carbon storage capacity and GHG emission characteristics (Sasmith et al., 2019). Such disturbances may lead to a regime shift in the ecosystem's carbon sink pattern, e.g., the disturbances transformed the ecosystem from a stable CO<sub>2</sub> sink into a weak source (Zhu et al., 2024b). The removal and burial of cordgrass may lead to intensive decomposition and decay of residual biomass, potentially releasing substantial amounts of GHGs (Yang et al., 2024) with short-term pulse emission (Pendleton et al., 2012). However, there is still a lack of accurate quantification and estimation of GHG fluxes via direct empirical data, especially the CH<sub>4</sub> fluxes before and after the cordgrass removal.

In BCEs, CH<sub>4</sub> emission primarily originates from the microbial decomposition of soil organic matter, with the availability of organic substrates often being a key limiting factor for CH<sub>4</sub> production (Al-Haj and Fulweiler, 2020; Noyce and Megonigal, 2021; Hu et al., 2024). Notably, anthropogenic disturbances can trigger pulse GHG emission, typically associated with the short-term input of large amounts of organic matter and the rapid decomposition of labile substrates. For example, it has been shown that the mortality of belowground biomass in disturbed mangroves provides abundant substrates for microbial decomposition, significantly enhancing heterotrophic respiration and substantially increasing CO<sub>2</sub> emission (Sasmith et al., 2019). The ultimate CH<sub>4</sub> efflux entering the atmosphere is the result of multiple biogeochemical processes, including CH<sub>4</sub> production, oxidation, and transport (Bridgman et al., 2013). These complex processes are jointly regulated by various environmental factors, among which temperature plays a particularly prominent role. It is widely recognized that temperature directly and indirectly controls soil biogeochemical processes by modulating microbial metabolism and influencing substrate availability (Wang et al., 2015). Higher temperature not only significantly enhances the metabolic activity of methanogens but also accelerates organic matter decomposition, thereby providing more abundant substrates for CH<sub>4</sub> production (Hu et al., 2020). Emerging research further reveals that the temperature sensitivity of CH<sub>4</sub> emission may be closely linked to the decomposability of carbon substrates (Hu et al., 2024). Tidal activity represents another critical physical process influencing CH<sub>4</sub> fluxes. On the one hand, the changes in water table caused by tidal activities can affect CH<sub>4</sub> production and oxidation via altering oxygen availability (Hu et al., 2024), or modulate CH<sub>4</sub> release through hydrostatic pressure adjustments and tidal pumping effects (Wei et al., 2020; Cui et al., 2024; Zhu et al., 2024a). On the other hand, lateral tidal flows may introduce allochthonous organic carbon, promoting CH<sub>4</sub> production (Cui et al., 2024), while simultaneously facilitating the lateral export of sequestered blue carbon (Wang et al., 2023a). These intricate regulatory mechanisms with the interactions among various factors may amplify anthropogenic impacts on CH<sub>4</sub> fluxes, resulting in significant temporal variability in CH<sub>4</sub> emission from BCEs.

The changes in CH<sub>4</sub> flux dynamics induced by anthropogenic

disturbances from coastal restorations remain a critical knowledge gap in blue carbon science (Macreadie et al., 2021). To date, the degree to which CH<sub>4</sub> emission and its environmental controls changes with disturbance in BCEs has been rarely assessed with direct empirical evidence. A major constraint is the lack of sufficient CH<sub>4</sub> flux measurements with a high temporal resolution. The application of eddy covariance (EC) technique, capable of measuring high-frequency and continuous GHG fluxes between ecosystems and the atmosphere (Baldocchi, 2020; Zhang et al., 2022; Zhang et al., 2023; Deng et al., 2025), could serve as an efficient approach to help close this knowledge gap.

Here, to quantify these disturbance effects on CH<sub>4</sub> flux dynamics, we utilized the EC approach to continuously measure CH<sub>4</sub> flux from Jul. 2022 to Oct. 2023, covering both pre- and post-removal periods, in a disturbed saltmarsh-mangrove wetland of Southeast China experiencing an intensive cordgrass removal in late Oct. 2022. Specific objectives are (1) to quantitatively compare the differences in the temporal variability of CH<sub>4</sub> fluxes before and after the cordgrass removal, and (2) to assess the potential change in the response of CH<sub>4</sub> fluxes to key environmental controls including tidal activities and temperature dependence. We hypothesize that coastal restoration via intensive smooth cordgrass removal strengthens tidal and temperature impacts and thus promotes the magnitude of CH<sub>4</sub> emission and its diel variability (Fig. 1). This study should offer direct empirical evidence to improve the understanding of the effects of restoration-induced anthropogenic disturbances on the temporal variability and environmental controls of CH<sub>4</sub> emission.

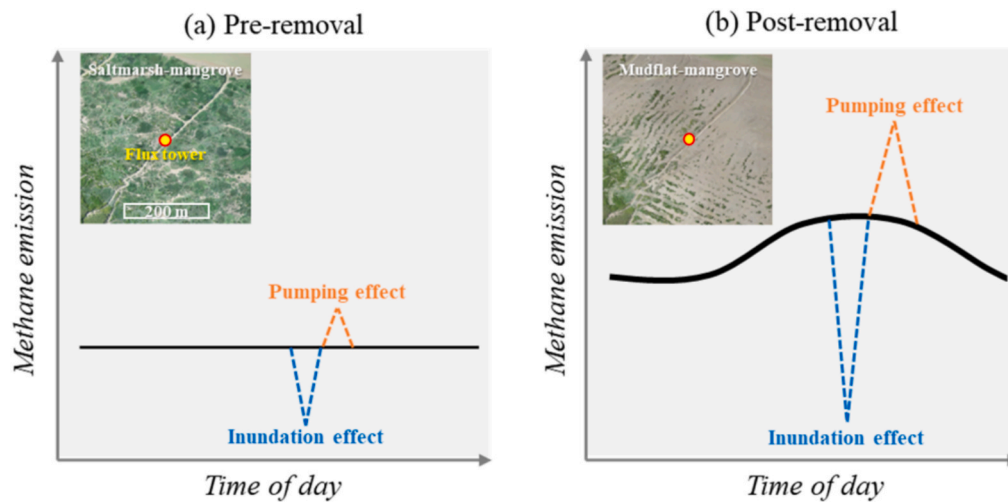
## 2. Materials and methods

### 2.1. Study area

A flux tower (23.9201°N, 117.4227°E; part of ChinaFLUX and USCCC) was established in a disturbed tidal wetland experiencing cordgrass removals, within Zhangjiang Mangrove National Nature Reserve in Southeast China, to conduct EC measurements of net CH<sub>4</sub> exchange (NME) between the wetland and the atmosphere (Zhu et al., 2021c). This wetland has a subtropical monsoon climate, with a mean air temperature of 21.1°C and annual rainfall of 1165.5 mm, and experiences a semidiurnal tidal regime, with a mean tidal range of 2.1 m and surface water salinity of 12.6 ppt (Lu and Zhu, 2021; Zhu et al., 2019a; Zhu et al., 2021b). This wetland has experienced range expansions of both invasive cordgrass and native mangroves over the past decade to form a saltmarsh-mangrove ecotone (Fig. 1a) (Zhu et al., 2019b; Zhu et al., 2024b). Since late Oct. 2022, the government initiated a large-scale coastal wetland restoration effort to intensively remove cordgrass and thus turned the ecotone into a mudflat-mangrove habitat (Fig. 1b) within a few weeks (Zhu et al., 2025). Successive removal campaigns were maintained to prevent the cordgrass regrowth, including a moderate one occurring in late Jun. 2023. According to the detailed removal activities revealed by in-situ camera monitoring (Zhu et al., 2024b), we defined a pre-removal period from Jul. 8th, 2022 to Oct. 20th, 2022 and a post-removal period from Nov. 1st, 2022 to Oct. 31st, 2023.

### 2.2. Field measurements

The 30-min NME time series data were derived from raw 10-Hz measurements by an open-path EC system of a three-axis anemometer (CSAT-3, Campbell Scientific, Inc., Logan, UT, USA) and a CH<sub>4</sub> gas analyzer (Li-7700, Li-COR Inc., Lincoln, NE, USA). The 10-Hz data were processed into 30-min data following a series of flux correction (e.g., axis rotation, frequency response, ultrasonic, and WPL corrections) and quality control (e.g., steady state, turbulent condition, statistical, and absolute limit tests) processes, as implemented in the EddyPro software (Li-COR Inc.) (Zhu et al., 2019c; Zhu et al., 2021a; Liu and Zhu, 2024). The flux data quality was labeled using the 0–1–2 flagging system (Mauder et al., 2013), and the data with a quality flag of 2 were excluded



**Fig. 1.** Conceptual diagram illustrating the changes in methane emission and its diel variability along with the smooth cordgrass removal in a disturbed tidal wetland. Wetland landscape around the flux tower during the (a) pre-removal (saltmarsh-mangrove habitat) and (b) post-removal (mudflat-mangrove habitat) periods are shown by two insets of drone mosaic images acquired in Jun. 2022 and Jul. 2023, respectively. Tidal inundation and pumping are expected to suppress and promote methane emission, respectively.

from further analyses (Zhu et al., 2021a; Wang and Zhu, 2024). Additionally, flux data under conditions of rainfall and insufficient nighttime turbulence were also excluded (Zhu et al., 2021c). Regular system maintenance (e.g., manually cleaning Li-7700 mirrors) was conducted to ensure the data quality and the NME data with signal strength below 20% were excluded (Zhu et al., 2021c). Footprint climatology analyses confirmed that 80% of the fluxes were contributed from the habitats within  $\sim 200$  m around the tower, dominated by southeast and northwest wind directions (Zhu et al., 2024b). As a result of quality controls and system malfunctions, the coverage of valid 30-min NME data was 62.9% (66.2% for daytime and 59.5% for nighttime) over the study period from Jul. 8th, 2022 to Oct. 31st, 2023. Auxiliary 30-min meteorological and tidal data included photosynthetically active radiation (PAR; PQS1 PAR Quantum sensor, Kipp & Zonen, Delft, Netherlands), air temperature (Tair; HMP155A sensor, Vaisala, Helsinki, Finland), rainfall (TE525MM Rain Gage, Campbell Scientific, Inc.), and surface water level (HOBO U20L-04 Water Level Logger, Onset, Bourne, MA, USA).

### 2.3. Data analyses

The 30-min NME data were temporally aggregated to daily data. To ensure a minimum representative coverage of the diel cycle while retaining a sufficient number of days for analysis, the days with daytime valid 30-min records below one-fourth or nighttime valid records below one-eighth were excluded from the temporal aggregation. A solar elevation angle threshold of zero was applied to distinguish daytime and nighttime periods. To reduce the effect of the uneven data missing issue, the mean NME was calculated separately for daytime and nighttime and then temporally aggregated to daily data. Daily data were further averaged to monthly data. Over the studying period, the percentage of valid daily NME was 77.3%. To examine the difference in NME between daytime and nighttime, we used an independent *t*-test analysis to statistically compare daytime and nighttime 30-min NME for each month over the studying period. Pearson correlation analyses were used to examine the correlations among 30-min NME and relevant environmental variables for both pre- and post-removal periods. To examine the tidal impacts on NME, mean diel variation (MDV) analyses were conducted to compare the diel variability of NME under tidal inundated and exposed conditions. To explore the correlation between air temperature and NME, Pearson correlation coefficients were also calculated between 30-min NME and air temperature for each single day. To better examine

the change in the temperature dependence of NME over time, we quantified the temperature sensitivity of  $Q_{10}$  for each month. The NME data acquired under tidal exposed conditions were used only to avoid the confounding effect of tidal inundation. Specifically, all 30-min NME data under tidal exposed conditions for each month were bin averaged into 10 groups by the percentiles of air temperature and then the exponential equation with parameters *a* and *b* (Eq. (1); Wang et al., 2015) was used to fit the empirical relationship between temperature percentiles (*T*) and binned median NME. Thus, monthly  $Q_{10}$  was calculated from the fitted parameter *b* (Eq. (2)).

$$NME = a \cdot \exp^{b \cdot T} \quad (1)$$

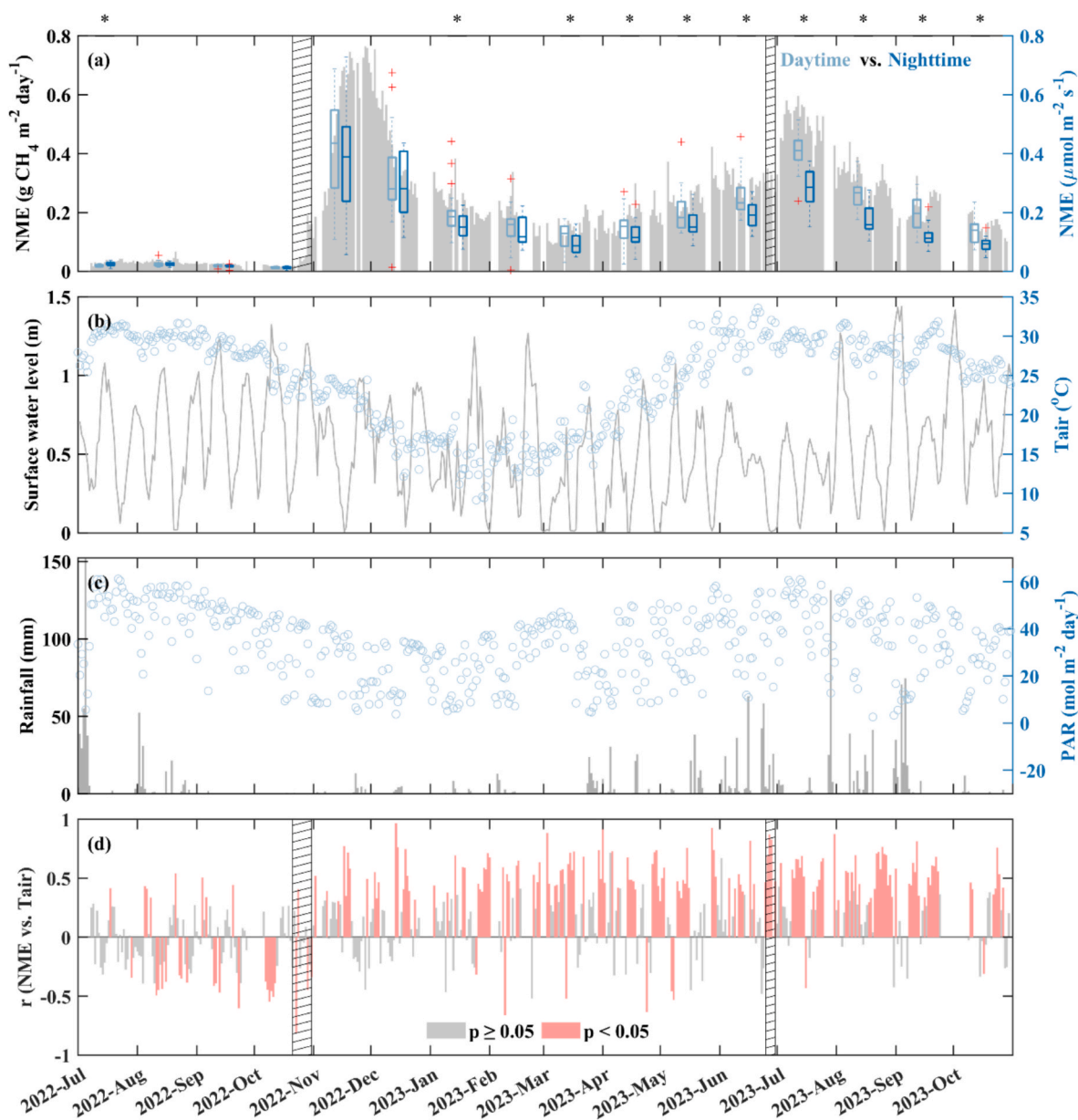
$$Q_{10} = \exp^{10b} \quad (2)$$

In this study, a positive (negative) NME value indicated an upward (downward)  $CH_4$  flux or a  $CH_4$  source (sink). All data processing and statistical analyses were performed in MATLAB software (MathWorks Inc., Natick, MA, USA).

## 3. Results

### 3.1. Temporal variations in methane emission

Over the pre-removal period, daily NME varied from  $0.008 \text{ g } CH_4 \text{ m}^{-2} \text{ d}^{-1}$  to  $0.067 \text{ g } CH_4 \text{ m}^{-2} \text{ d}^{-1}$  and monthly mean NME changed from  $0.018 \pm 0.003$  (mean  $\pm$  std.)  $\text{g } CH_4 \text{ m}^{-2} \text{ d}^{-1}$  in Oct. 2022 (excluding the end of month during the removal period) to  $0.034 \pm 0.009 \text{ g } CH_4 \text{ m}^{-2} \text{ d}^{-1}$  in Aug. 2022 (Fig. 2a). The removals significantly boosted the emission with pulse increases following each of the two major removals (intensive and moderate ones). The peak emission of the first pulse approximately one month later ( $0.764 \text{ g } CH_4 \text{ m}^{-2} \text{ d}^{-1}$ ) was 27.3 times that of the pre-removal level ( $0.028 \pm 0.008 \text{ g } CH_4 \text{ m}^{-2} \text{ d}^{-1}$ ), while the peak emission of the second pulse ( $0.596 \text{ g } CH_4 \text{ m}^{-2} \text{ d}^{-1}$ ) was twice that of the previous month ( $0.301 \pm 0.034 \text{ g } CH_4 \text{ m}^{-2} \text{ d}^{-1}$ ). Over the post-removal period, the two highest mean monthly NME occurred in the month right after each of the two major removals ( $0.530 \text{ g } CH_4 \text{ m}^{-2} \text{ d}^{-1}$  in Nov. 2022 and  $0.496 \text{ g } CH_4 \text{ m}^{-2} \text{ d}^{-1}$  in Jul. 2023), while the lowest mean monthly NME of  $0.146 \text{ g } CH_4 \text{ m}^{-2} \text{ d}^{-1}$  occurred in Mar. 2023. Mean daily NME over the full post-removal period ( $0.297 \pm 0.155 \text{ g } CH_4 \text{ m}^{-2} \text{ d}^{-1}$ ) was 10.6 times that of the pre-removal level. The direct comparison of NME for the same time window (Jul. 8th  $\sim$  Oct. 20th) with available measurements both before and after the removal indicated that the post-removal mean



**Fig. 2.** Temporal variability of methane emission and environmental factors from Jul. 2022 to Oct. 2023 in a disturbed tidal wetland experiencing smooth cordgrass removals, including (a) daily/monthly net methane exchange (NME), (b) daily maximum surface water level, daily mean air temperature ( $T_{\text{air}}$ ), (c) daily rainfall, and daily photosynthetically active radiation (PAR). The differences in 30-min NME between daytime (light blue) and nighttime (dark blue) periods for each month are shown by a pair of boxplots, with an asterisk above the upper x-axis indicating a statistical significance ( $p < 0.05$ ) (a). Daily Pearson correlation coefficients ( $r$ ) between 30-min NME and  $T_{\text{air}}$  are shown with statistically significant (red) and non-significant (gray) values (d). The two diagonal shadings indicate the intensive (late Oct. 2022) and moderate (late Jun. 2023) removals, respectively (a and d).

daily NME ( $0.317 \pm 0.128 \text{ g CH}_4 \text{ m}^{-2} \text{ d}^{-1}$ ) was 11.3 times higher than the pre-removal level (Fig S1). Overall,  $\text{CH}_4$  emission increased significantly following the intensive removal, sustaining a pulse emission for approximately two months before a winter decline. Thereafter, emission rebounded from the beginning of the growing season (Mar. 2023) onward until the moderate removal event triggering another emission pulse, after which a sustained decreasing trend was observed.

Over the pre-removal period, no statistically significant difference ( $t$ -test) was observed between daytime and nighttime NME except Jul. 2022 (Fig. 2a). Monthly-mean daytime NME ranged from  $0.018 \pm 0.004 \mu\text{mol m}^{-2} \text{ s}^{-1}$  in Sep. 2022 to  $0.024 \pm 0.008 \mu\text{mol m}^{-2} \text{ s}^{-1}$  in Aug. 2022, while nighttime NME varied between  $0.012 \pm 0.004 \mu\text{mol m}^{-2} \text{ s}^{-1}$  in Oct. 2022 (excluding the end of month during the removal period) and  $0.025 \pm 0.006 \mu\text{mol m}^{-2} \text{ s}^{-1}$  in Aug. 2022, with a diel difference (i.

e., daytime minus nighttime) of  $-0.004$ – $0.0008 \mu\text{mol m}^{-2} \text{ s}^{-1}$ . In the month following the intensive removal (Nov. 2022), mean daytime and nighttime NME increased to  $0.426 \pm 0.170 \mu\text{mol m}^{-2} \text{ s}^{-1}$  (34.5 times that of Oct. 2022) and  $0.379 \pm 0.175 \mu\text{mol m}^{-2} \text{ s}^{-1}$  (31.2 times), respectively. Over the post-removal period, the mean daytime NME of  $0.237 \pm 0.131 \mu\text{mol m}^{-2} \text{ s}^{-1}$  (nighttime:  $0.187 \pm 0.112 \mu\text{mol m}^{-2} \text{ s}^{-1}$ ) was 11.9 (8.9) times that of the pre-removal daytime level of  $0.020 \pm 0.007 \mu\text{mol m}^{-2} \text{ s}^{-1}$  ( $0.021 \pm 0.007 \mu\text{mol m}^{-2} \text{ s}^{-1}$ ). The lowest post-removal NME in Mar. 2023 (daytime:  $0.118 \pm 0.041 \mu\text{mol m}^{-2} \text{ s}^{-1}$ ; nighttime:  $0.091 \pm 0.031 \mu\text{mol m}^{-2} \text{ s}^{-1}$ ) were still 5.9 and 4.3 times that of the pre-removal levels, respectively.  $T$ -test statistical analyses confirmed that daytime NME was statistically significantly larger than nighttime NME in most months over the post-removal period. These diel differences were further revealed by the monthly MDV analyses (Fig. 3).



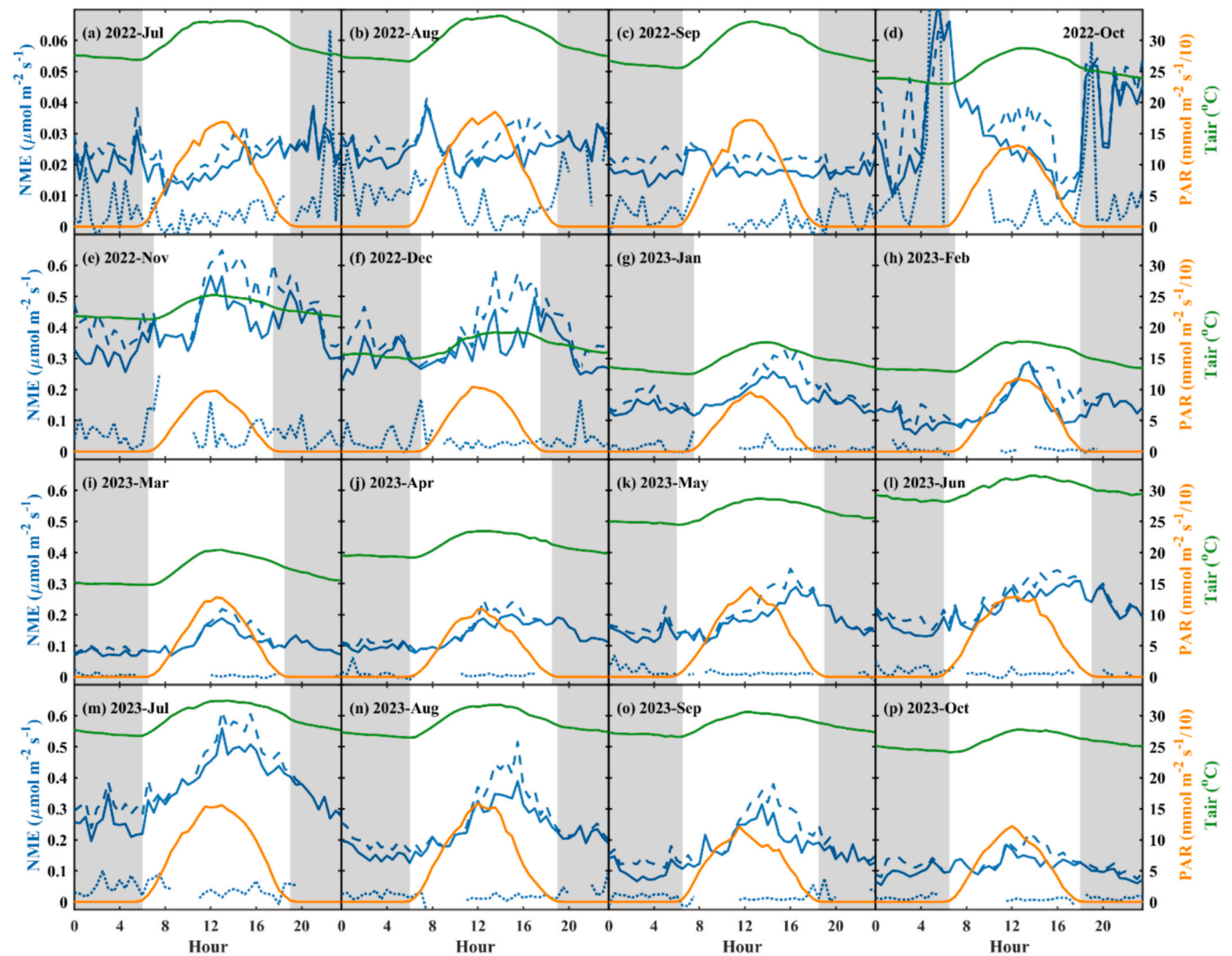


Fig. 3. Mean diel variations in 30-min net methane exchange (NME; blue), photosynthetically active radiation (PAR; orange), and air temperature (Tair; green) for each month from Jul. 2022 to Oct. 2023. Mean diel variations in 30-min NME under tidal exposed (dashed blue) and inundated (dotted blue) conditions are also shown for each month. Note that the scale of NME used for the first row was different from that for the other rows.

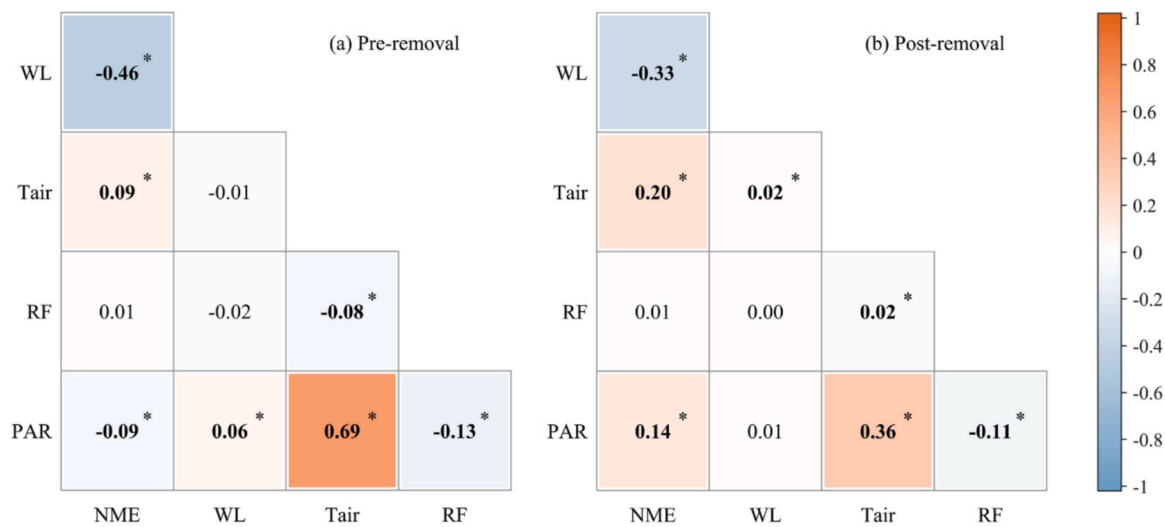


Fig. 4. Heatmaps of Pearson correlation coefficients during the pre-removal (a) and post-removal (b) periods among 30-min net methane exchange (NME) and environmental factors, including surface water level (WL), air temperature (Tair), rainfall (RF), and photosynthetically active radiation (PAR). Statistically significant ( $p < 0.05$ ) coefficients are in bold and marked with asterisks.

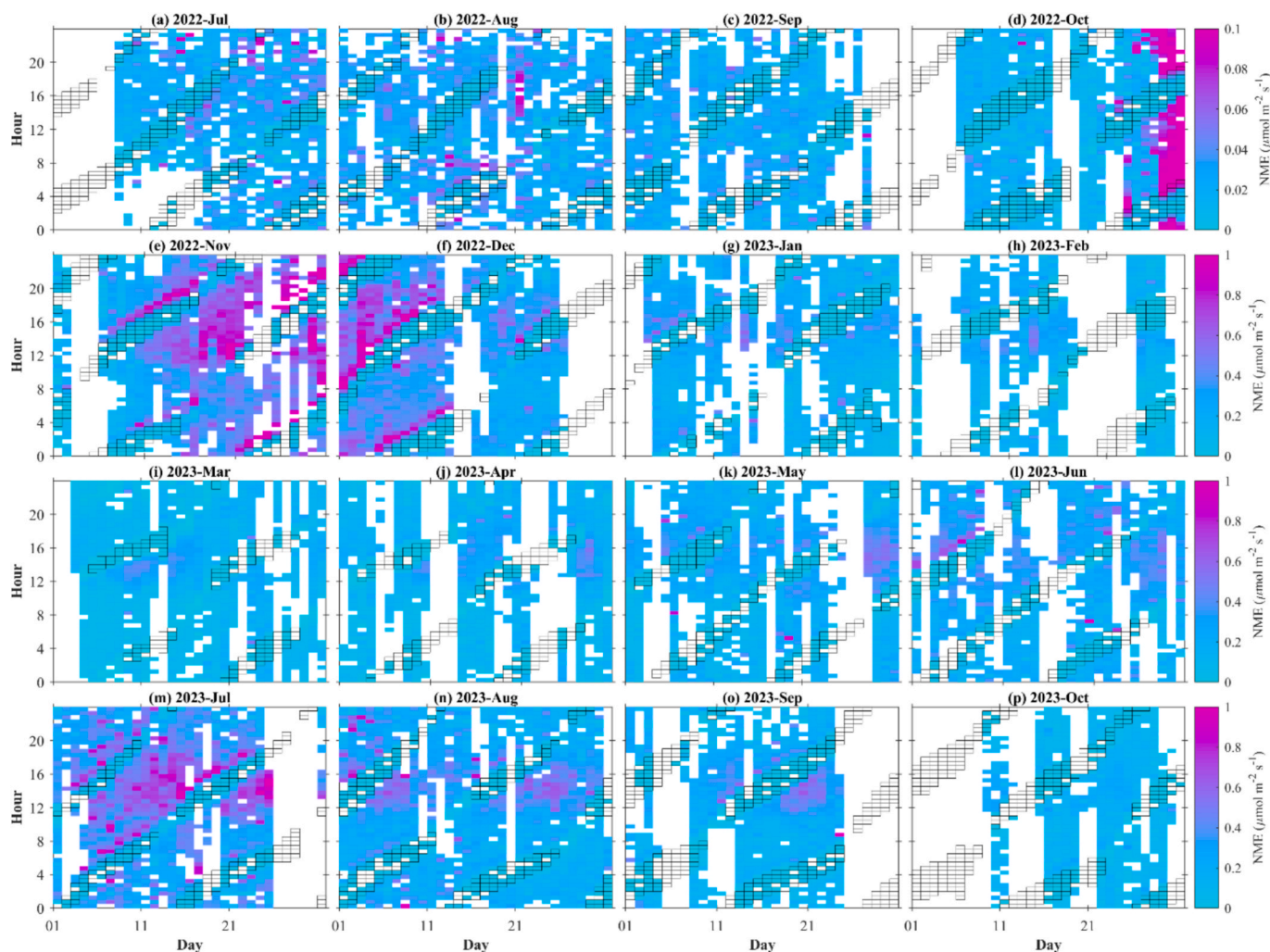
Over the pre-removal period, NME was relatively constant over the course of a day without obvious diel differences, while over the post-removal period NME had much stronger diel variability with higher daytime emission occurring in the afternoon.

### 3.2. Environmental controls on methane emission

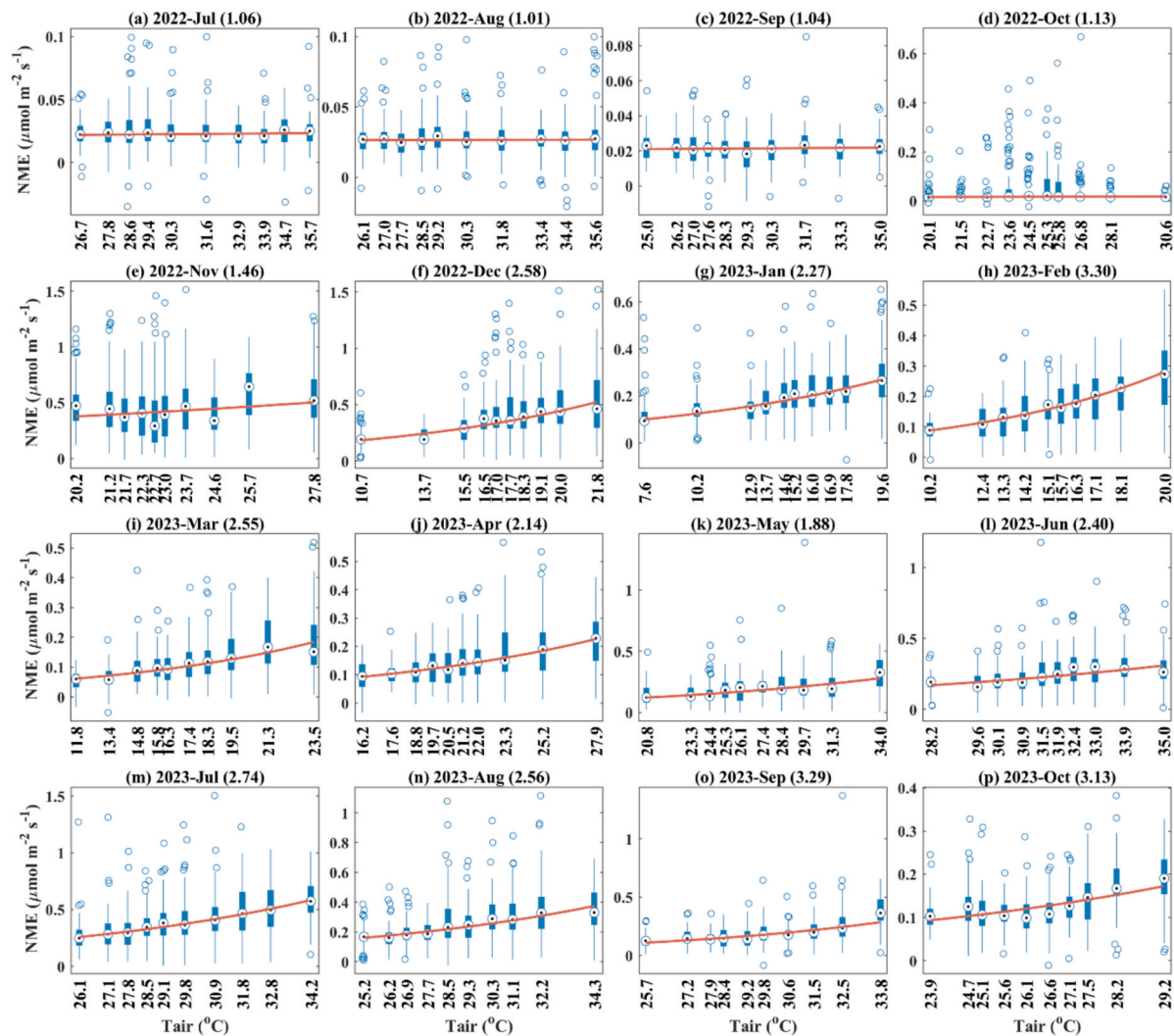
Potential environmental controls on  $\text{CH}_4$  emission were explored using Pearson correlation analyses between 30-min NME and environmental factors (Fig. 4). Over the pre-removal period, 30-min NME was statistically significantly ( $p < 0.05$ ) correlated with WL (correlation coefficient:  $-0.46$ ), Tair ( $0.09$ ) and PAR ( $-0.09$ ), while its correlation with RF ( $0.01$ ) was not statistically significant. Over the post-removal period, 30-min NME was also statistically significantly correlated with WL ( $-0.33$ ), Tair ( $0.20$ ) and PAR ( $0.14$ ). Given that both pre- and post-removal NME were statistically significantly correlated with WL and Tair on a 30-min scale, we further examined tidal and temperature impacts on NME on a sub-daily scale. Based on the monthly MDV analyses,  $\text{CH}_4$  emission under inundated conditions was obviously lower than that under exposed conditions (Fig. 3). Over the pre-removal period, mean NME under inundated conditions ranged from  $0.003$  to  $0.013 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ , while mean NME under exposed conditions was  $2.6$ – $6.5$  times higher, ranging from  $0.022$  to  $0.037 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ . Over the post-removal period, mean NME under inundated conditions varied

from  $0.006$  to  $0.059 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ , while mean NME under exposed conditions was  $9.0$ – $27.0$  times higher, ranging from  $0.114$  to  $0.456 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ . Based on the 30-min time series of NME, pronounced tidal pumping effects were well identified, in which the efflux surged abruptly at the transition from inundated to exposed conditions (Fig. 5). This pumping effect was obviously stronger during the two months following the removal events. Direct comparison of 30-min NME before and after the inundation-exposure transition indicated that, over the pre-removal period, the efflux right after the transition was  $2.0$ – $3.4$  times higher, while this difference increased to  $4.6$ – $9.9$  times over the post-removal period.

Although pulse emissions occurred after the removal, the seasonal variability of  $\text{CH}_4$  efflux underlying the pulse emissions still matched the seasonal pattern of Tair (Fig. 2a–b). Nearly half (48.9%) of valid days showed statistically significant correlations between 30-min NME and Ta over the post-removal period, while there was only 27.1% over the pre-removal period (Fig. 2d). Pearson correlation coefficients over the post-removal period were overall higher than those over the pre-removal period. Over the pre-removal period, weak temperature dependence of NME was observed with monthly  $Q_{10}$  close to one ( $1.01$ – $1.13$ ), while over the post-removal period, the temperature dependence obviously increased with  $Q_{10}$  varying from  $1.46$  in Nov. 2022 to  $3.30$  in Feb. 2023 (Fig. 6). On average, the post-removal  $Q_{10}$  values were  $2.39$  times that of the pre-removal level. By comparing the



**Fig. 5.** Temporal variability of 30-min net methane exchange (NME) in a disturbed tidal wetland experiencing smooth cordgrass removals for each month from Jul. 2022 to Oct. 2023. The grids with a black outline indicate those 30-min periods experiencing tidal inundation. Note that the rightmost color bar applies to the four months in the same row, with the first row using different scales from the others for better visibility.



**Fig. 6.** Response of net methane exchange (NME) to air temperature ( $T_{air}$ ) under tidal exposed conditions from Jul. 2022 to Oct. 2023 in a disturbed tidal wetland experiencing smooth cordgrass removals. All 30-min NME data acquired under exposed conditions for each month are bin-averaged into 10 groups by the percentiles of  $T_{air}$  in that month, and then an exponential response function is applied to fit the binned median NME and corresponding percentiles of  $T_{air}$ , as shown on each X-axis. The number in brackets in the sub-title denotes the  $Q_{10}$  value of each month. Note that both the X-axis and Y-axis for each month use different scales.

same months (Jul., Aug., and Sep.) between pre- and post-removal periods, the post-removal  $Q_{10}$  values were 2.53–3.16 times that of the pre-removal levels. Over the studying period, there was an overall statistically positive relationship between  $Q_{10}$  and NME on a monthly scale ( $y = 0.026\exp^{0.83x}$ ,  $R^2 = 0.44$ ) (Fig. 7).

## 4. Discussion

### 4.1. Changing methane emission with the removal

Previous studies have found that the cordgrass invasion enhances  $CH_4$  emission by altering vegetation productivity, soil properties, and microbial community structure, as well as the characteristics of plant-mediated  $CH_4$  transport, thereby profoundly impacting the soil carbon cycling in coastal wetlands (Yuan et al., 2015; Bu et al., 2019; Liu et al., 2022; Zheng et al., 2023). Similarly, we find the  $CH_4$  fluxes observed in this saltmarsh with invasive cordgrass, ranging from 466.9 to 4203.7  $\mu\text{mol } CH_4 \text{ m}^{-2} \text{ d}^{-1}$  before the removal, are significantly higher than the average value of the saltmarsh (224.4  $\mu\text{mol } CH_4 \text{ m}^{-2} \text{ d}^{-1}$ ) reported by a synthesis study (Al-Haj and Fulweiler, 2020). Given that mowing combined with mechanical plowing is used in the cordgrass removal, the pulse  $CH_4$  emission following the removals can be explained by several

aspects. First, the large amount of fragmented plant residues buried by plowing provides abundant fresh carbon substrates for methanogens (Yang et al., 2021), which is similar to the effect of straw incorporation (Zhang et al., 2015). Second, the mechanical soil disturbance and residue decomposition alter microbial community composition (Wang et al., 2023b), favoring fast-growing decomposers that accelerate organic matter breakdown and methanogenic activity (Li et al., 2022). Third, the soil compaction from heavy machinery can form a more stable and anoxic environment that greatly facilitates the methanogenesis (Villa, 2020).

There is a clear positive correlation between the labile organic matter and  $CH_4$  production (Bastviken et al., 2023). Studies have shown that the litter of cordgrass primarily consists of labile organic matter that is easily decomposed (Li et al., 2024; Huang et al., 2025). Mowing and plowing result in a substantial short-term input of the labile organic matter into the soil and thus greatly stimulate  $CH_4$  production and emission. In this study,  $CH_4$  efflux is maintained at high levels for two months after the cordgrass removal, and then gradually decline and stabilize likely due to the exhaustion of the pulse input of labile organic matter (Fig. 2). Although the peak value of second pulse emission in Jul. 2023 is lower than that of first pulse emission, the  $CH_4$  fluxes following the moderate removal remain markedly higher than the same months of



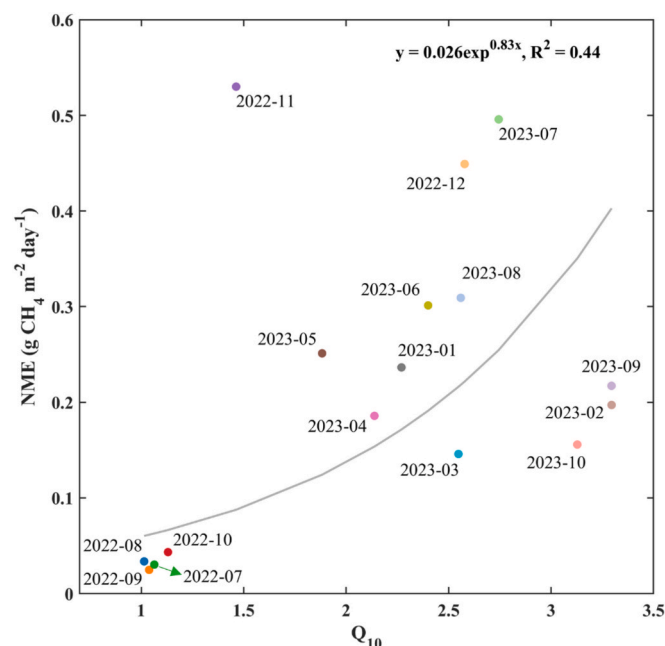


Fig. 7. Response of monthly mean net methane exchange (NME) to its temperature sensitivity ( $Q_{10}$ ) from Jul. 2022 to Oct. 2023 in a disturbed tidal wetland experiencing smooth cordgrass removals, shown with a fitted exponential response curve.

the previous year. These findings suggest that the substrate availability may be the key determinant of the magnitude and duration of CH<sub>4</sub> pulses. The finding of large changes in the magnitude and temporal variability of CH<sub>4</sub> emission implies that these anthropogenic disturbances may potentially lead to a regime shift in soil CH<sub>4</sub> biogeochemistry. Given that this restoration also causes a large change in CO<sub>2</sub> flux (Zhu et al., 2024b), we also conduct a comparative analysis, using the data within the same time window (Jul. 8th ~ Oct. 20th) before and after the removal, to further examine how the combined effects on both CH<sub>4</sub> and CO<sub>2</sub> fluxes determine the net climate benefit (Fig S1). The comparison reveals that the estimated net climate benefit, based on the sustained-flux global warming potentials metric for a 100-year time horizon (Zhu et al., 2025), changes from a pre-removal cooling effect ( $-12.03 \pm 6.36$  g CO<sub>2</sub> eq. m<sup>-1</sup> day<sup>-1</sup>) to a post-removal warming effect ( $15.26 \pm 6.23$  g CO<sub>2</sub> eq. m<sup>-1</sup> day<sup>-1</sup>). It is evident that the combination of the CH<sub>4</sub> pulse emission and the reversal of CO<sub>2</sub> flux leads to a significant negative climate benefit.

#### 4.2. Enlarged tidal impacts on methane emission

Tidal activity is an important factor influencing carbon exchange in BCEs. As a result of the suppressing effects of tidal inundation, the CH<sub>4</sub> fluxes during tidal inundation in this study are significantly lower than those during tidal exposure (Fig. 3). On the one hand, tidal inundation alters soil redox potential and creates a suitable anaerobic environment conducive to CH<sub>4</sub> production while simultaneously constraining CH<sub>4</sub> oxidation. Tidal inundation may also introduce more sediment inputs, indirectly enhancing methane generation (Cui et al., 2024). On the other hand, a large portion of the produced CH<sub>4</sub> diffuses into porewater or overlying water (Hu et al., 2020). Tidal inundation just acts as a physical barrier and inhibiting the release of CH<sub>4</sub> to the atmosphere (Wei et al., 2020), leading to short-term CH<sub>4</sub> accumulation during tidal inundation and pulse emission during tidal exposure (Fig. 5). Additionally, due to tidally driven pore water exchange, the dissolved carbon released from the cordgrass residues are transported laterally into adjacent tidal creeks (Capooci et al., 2024; Huang et al., 2025), where they may be remineralized into CH<sub>4</sub> and emitted into the atmosphere within a short period

of time (Bogard et al., 2020). This is likely an additional contributor to the pulse CH<sub>4</sub> emission (Rosentreter et al., 2018).

Tidal pumping is a tidally driven exchange process between porewater and surface water, which significantly influences GHG emissions at the water–air interface in BCEs (Call et al., 2015; Chen et al., 2021a). Such CH<sub>4</sub> emission hotspots induced by tidal pumping have been widely reported in various BCEs, including mangroves and salt marshes (Li et al., 2018; Rosentreter et al., 2018; Zhu et al., 2024a). Over the post-removal period, with little plant-mediated transport through aerenchyma, CH<sub>4</sub> emission primarily occurs via diffusion and ebullition. During ebb tide, the reduction in hydrostatic pressure facilitates the release of CH<sub>4</sub> supersaturated in porewater (Rosenqvist et al., 2002). In the absence of plant-mediated transport, CH<sub>4</sub> accumulated in porewater tends to release rapidly as bubbles, minimizing the oxidative loss of CH<sub>4</sub> during the transport process (Bastviken et al., 2023). Thus, ebullition may be the dominant pathway for CH<sub>4</sub> pulse emission following cordgrass removal. Although the aerenchyma of cordgrass plays a crucial role in CH<sub>4</sub> transport (Comer-Warner et al., 2022), the contribution of ebullition cannot be overlooked, especially under conditions of abundant organic matter (e.g., decaying plant residues) (Bastviken et al., 2023). Notably, tidal inundation and pumping effects occur both before and after the removal, but the difference in CH<sub>4</sub> fluxes between inundated and exposed conditions is significantly greater during the post-removal period. These findings suggest that the removals strengthen tidal controls on CH<sub>4</sub> emission and its diel variability by enhanced tidal inundation and priming effects.

#### 4.3. Enhanced temperature sensitivity of methane emission

It is assumed in this study that the change in organic substrates after the removal is a key factor contributing to the enhanced temperature sensitivity. The  $Q_{10}$  value could be affected by the quality and quantity of organic substrates. This explanation is consistent with previous studies that emphasize the critical role of soil substrate availability in modulating the temperature sensitivity of CH<sub>4</sub> emission. An abrupt increase in substrate availability can lead to a higher apparent temperature sensitivity. For instance, Chang et al. (2019) demonstrate that increased labile substrates following peatland thawing heighten CH<sub>4</sub> temperature sensitivity, and Wang et al. (2021) indicate that typhoon and rainstorm events deliver more terrestrial organic matter to wetlands and thus stimulate CH<sub>4</sub> emission. Water table fluctuations in wetland can also modulate the temperature sensitivity by regulating the substrate accessibility and microbial activity (Chen et al., 2021b; Li et al., 2023; Hu et al., 2024). Moreover, experimental evidence reveals that increased carbon availability exerts more positive impacts on the temperature sensitivity of methanogenesis compared to CH<sub>4</sub> oxidation (Inglett et al., 2012). It is possible that the pulse of nutrients released during the decay of cordgrass promotes the establishment of methanogenesis-related microbial communities (Wang et al., 2023b), which not only accelerate methanogenesis but also contribute to a more sensitive temperature response of CH<sub>4</sub> production.

Another possible reason for the enhanced temperature sensitivity is that the cordgrass removal reduces the oxygen availability and thus CH<sub>4</sub> oxidation, leading to a net increase in CH<sub>4</sub> emission. Over the pre-removal period, the diel variation in CH<sub>4</sub> emission is relatively small (Figs. 2a and 3), which is in contrast with the CH<sub>4</sub> flux patterns observed in adjacent mangrove ecosystems (Zhu et al., 2024a). This is likely explained by the fact that daytime CH<sub>4</sub> fluxes are suppressed due to enhanced rhizosphere CH<sub>4</sub> oxidation driven by plant-mediated oxygen transport (Noyce et al., 2023). This explanation is supported by Pearson correlation analyses, showing that the relationship between CH<sub>4</sub> fluxes and PAR reversed from a significant negative correlation before the removal to a significant positive correlation after the removal (Fig. 4). Higher daytime temperature promotes both CH<sub>4</sub> production and oxidation, and thus there is a possibility of no distinct diel variations in net CH<sub>4</sub> emission over the pre-removal period. However, over the post-



removal period, the proportion of daytime CH<sub>4</sub> oxidation could decrease due to lower oxygen supply to the rhizosphere, which leads to an increase in the diel variability of CH<sub>4</sub> emission since higher daytime temperature can significantly stimulate the substrate-rich methanogenesis (Knox et al., 2021).

#### 4.4. Limitations and implications

This study employs the EC technique to monitor the changes in CH<sub>4</sub> fluxes before and after the removal of cordgrass. However, several limitations remain in this study. First, although the EC technique provides high-frequency flux measurements, the data inevitably suffer from uncertainties due to the necessary quality controls and aggregation procedures. To reduce the uncertainty, we apply separate quality control for daytime and nighttime data before aggregating them to daily and monthly scales, which is more methodologically sound (Zhu and Zhu, 2025). Second, although the time-varying footprint coverage may influence the temporal dynamics of NME, our previous study has demonstrated that there is not much difference in footprint climatology during the study period with low coverage of newly established mangroves (Zhu et al., 2024b). Thus, the uncertainty from time-varying footprint coverage should be relatively small. Third, we use the air temperature as a surrogate for soil temperature due to the lack of soil temperature measurements in this study, which may introduce additional uncertainty due to the difference in magnitude and timing between air and soil temperature (Zhu et al., 2024a). Fourthly, we only apply a simple temperature-emission exponential function to examine the temperature sensitivity of CH<sub>4</sub> emission, despite that multiple factors, including temperature, microbial activity, and substrate availability actually influence the emission. Thus, empirical model parameters from these simple fits may not fully reflect the temperature sensitivity of observed flux dynamics (Wang et al., 2015). Fifthly, in addition to the removal activities, the difference in climate conditions between the pre- and post-removal years also affects the interannual variations in observed CH<sub>4</sub> fluxes. Following Zhu et al. (2024b), we partition the removal- and climate-induced contribution (Fig S2) and confirm that the change in CH<sub>4</sub> fluxes is predominantly attributed to the cordgrass removal instead of the climatic difference. Sixthly, although the high-frequency time series of observed CH<sub>4</sub> fluxes confirm that the cordgrass removals trigger CH<sub>4</sub> pulse emissions, our observations covering a one-year post-removal period are not long enough. Future longer flux measurements with extended years are needed to comprehensively assess the long-term evolution of CH<sub>4</sub> fluxes following the removals. Finally, given that the restoration-induced impacts on CH<sub>4</sub> emission depend on the disturbance types, vegetation composition, and environmental conditions, our findings may not be directly extrapolated to other restoration scenarios.

This study confirms the enhancements of both the magnitude and diel variability of CH<sub>4</sub> emission following the cordgrass removal. The disturbance-induced pulse emissions with high diel variability suggest that traditional chamber-based low-frequency measurements cannot well track the strong temporal dynamics of CH<sub>4</sub> flux following intensive cordgrass removals. This fully underscores the unique advantages of high-frequency and continuous EC measurements in assessing GHG flux changes under anthropogenic disturbances (Macreadie et al., 2021). Under the background of global warming, the increased temperature sensitivity of CH<sub>4</sub> emission post-removal implies that future climate warming may induce more intense CH<sub>4</sub> emission. Notably, cordgrass removal often requires repeated implementation, and the short-term CH<sub>4</sub> emission resulting from such periodic anthropogenic disturbances could create a low-frequency but high-magnitude amplification effect in future carbon budgets. Therefore, it is essential to consider GHG emission throughout the entire land-use change process in ecological restoration projects. This is because such CH<sub>4</sub> pulses of limited duration but considerable intensity pose a significant risk of offsetting the blue carbon benefits brought by ecological restoration.

## 5. Conclusions

The world's largest coastal ecosystem restoration is being implemented in coastal China to intensively remove invasive smooth cordgrass, potentially exerting a large impact on soil CH<sub>4</sub> emission. To assess the degree to which CH<sub>4</sub> emission and its environmental controls change with such anthropogenic disturbances, we utilize the eddy covariance method to track the temporal variability of CH<sub>4</sub> flux over both pre- and post-removal periods in a disturbed saltmarsh-mangrove wetland of Southeast China. Our high-frequency flux measurements reveal that the removal causes a pulse of CH<sub>4</sub> emission that is about one order of magnitude higher than the pre-removal level. We find that the removal not only intensifies the controls of tidal inundation and pumping on the CH<sub>4</sub> emission, but also enlarges the temperature sensitivity of the CH<sub>4</sub> emission. The combination of enhanced tidal impacts and temperature dependence thus promotes the diel variability of CH<sub>4</sub> flux over the post-removal period. These results highlight the necessity of better understanding the climate impact of restoration activities, which helps close the knowledge gap in assessing restoration-induced environmental risk of greenhouse gas flux changes. Future longer flux data with extended years are needed to further assess potential regime shift in soil CH<sub>4</sub> biogeochemistry and long-term evolution of such unintended environmental costs of the restoration.

#### CRedit authorship contribution statement

**Yueting Deng:** Writing – original draft, Visualization, Methodology, Investigation, Formal analysis, Data curation. **Ruichen Lin:** Visualization, Investigation, Data curation. **Han Yang:** Investigation, Data curation. **Hui Luo:** Investigation, Data curation. **Lulu Song:** Writing – review & editing, Supervision, Resources. **Xudong Zhu:** Writing – review & editing, Supervision, Resources, Project administration, Funding acquisition, Conceptualization.

#### Declaration of competing interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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#### Appendix A. Supplementary data

Supplementary data to this article can be found online at <https://doi.org/10.1016/j.geoderma.2026.117719>.

#### Data availability

Data will be made available on request.

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