

**RESEARCH ARTICLE**

**Methane emission intensifies the warming effect of carbon dioxide efflux from a subtropical coastal macroalgae aquaculture ecosystem**

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**Abstract**

Macroalgae aquaculture ecosystems have been increasingly recognized as coastal biogeochemical hotspots of air-sea net ecosystem carbon dioxide ( $\text{CO}_2$ ) exchange; however, their roles in regulating the temporal variability of net ecosystem methane ( $\text{CH}_4$ ) exchange (NME) receive little attention mainly due to very limited data availability. Here, we applied the eddy covariance (EC) technique to acquire 1-yr (June 2023 to May 2024) NME measurements, over a subtropical macroalgae aquaculture ecosystem in southeast China, to examine the temporal variability of NME across time scales and its contribution to net radiative forcing. The results indicated that (a) this ecosystem acted as a  $\text{CH}_4$  source in most months with the summer accounting for about two-thirds of annual NME of  $0.40 \text{ g C m}^{-2} \text{ yr}^{-1}$ ; (b) the inclusion of annual NME increased the sustained-flux global warming potentials (SGWPs) by 11.0% from  $219.3 \text{ g CO}_2\text{-eq. m}^{-2} \text{ yr}^{-1}$  for a 100-yr time horizon; (c) NME and its radiative contribution varied across seasons, farming periods, and growth stages, with the temporal fluctuations mainly controlled by temperature and tidal activities; (d) bimodal varying patterns across tidal levels were identified with larger fluxes occurring when tidal level changed most rapidly. This is the first EC study to confirm that  $\text{CH}_4$  emission intensifies the warming effect of  $\text{CO}_2$  efflux from macroalgae aquaculture ecosystems. The observed strong temporal variability of  $\text{CH}_4$  and  $\text{CO}_2$  fluxes and their asynchrony highlight the importance of high-frequency and continuous flux measurements in accurately assessing their net radiative forcing at both short- and long-term scales.

Macroalgae ecosystems have attracted more and more attentions in climate change mitigation because of their high primary production and carbon sequestration capacity compared to other coastal ecosystems (Krause-Jensen and Duarte 2016; Ortega et al. 2019; Queirós et al. 2019; Duarte et al. 2022). In addition to restoring natural macroalgae habitats, macroalgae aquaculture with sufficient cultivation area has recently garnered a lot of attention as a natural way to reduce greenhouse gases (GHGs) in the atmosphere (Bach et al. 2021; Gao et al. 2022). Macroalgae aquaculture is a

massive global industry, with China occupying the largest cultivation area of  $1252\text{--}1265 \text{ km}^2$  (Wu et al. 2020). Global macroalgae ecosystems have been found to sequester  $173 \text{ Tg C}$  of atmospheric carbon dioxide ( $\text{CO}_2$ ) annually (Krause-Jensen and Duarte 2016); however, methane ( $\text{CH}_4$ ) emission from these ecosystems has been often neglected (Al-Haj and Fulweiler 2020). Research has indicated that the coastal marine environment is a significant atmospheric  $\text{CH}_4$  source and actually contributes  $0.21 \text{ Pg CO}_2\text{-eq. yr}^{-1}$  to global  $\text{CH}_4$  emission (Rosentreter, Borges, et al. 2021; Resplandy et al. 2024). The availability of in situ  $\text{CH}_4$  flux measurements in coastal macroalgae ecosystems is very limited, which results in large uncertainties in the estimation of  $\text{CH}_4$  emission that hinders our efforts to promote the application

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of macroalgae aquaculture in terms of its role in climate change mitigation.

Coastal vegetated habitats provide sediment conditions that are favorable to high  $\text{CH}_4$  production because of the surplus organic matter resulting from extremely high primary production and the accumulation of allochthonous particulate organic matter (Ortega et al. 2019; Roth et al. 2022; Hall et al. 2025). Previous studies have shown that  $\text{CH}_4$  production in coastal waters primarily arises from methanogenesis in anoxic sediments (Ferrón et al. 2010; Hou et al. 2016). The  $\text{CH}_4$  consumption, occurring through anaerobic oxidation within sediments and aerobic oxidation within water columns (Knittel and Boetius 2009; Al-Haj and Fulweiler 2020), reduces the produced  $\text{CH}_4$  and thus net emission into the atmosphere at the air-sea interface. Diffusion, ebullition and plant-mediated transport are the three main pathways by which  $\text{CH}_4$  is released from sediments to the atmosphere in shallow coastal vegetated ecosystems (Jeffrey et al. 2019; Hilt et al. 2022). The former two emission pathways (i.e., diffusion and ebullition) are easily influenced by changing pressure conditions regulated by hydrodynamic factors such as tides and winds, leading to highly transient and pulsed  $\text{CH}_4$  flux on short timescales (Roth et al. 2022). Hence, pulse  $\text{CH}_4$  emission may occur from time to time given the high heterogeneity and variability of coastal environments. Overall,  $\text{CH}_4$  production, consumption, and transport involve complex processes that are modulated by a variety of direct and indirect factors, resulting in large temporal heterogeneity of  $\text{CH}_4$  flux (Rosentreter, Al-Haj, et al. 2021).

The  $\text{CH}_4$  flux and its contribution to the radiative forcing in macroalgae ecosystems may be inaccurately assessed due to a lack of enough in situ observations. On the one hand, coastal  $\text{CH}_4$  flux studies with high-frequency measurements focus on traditional blue carbon ecosystems including mangrove (J. Liu et al. 2020), saltmarsh (X. Zhu et al. 2025), and seagrass (Yau et al. 2023), while there are few such studies in macroalgae habitats, especially for macroalgae aquaculture. On the other hand, the exploration of the magnitude and temporal dynamics of  $\text{CH}_4$  flux may be impacted by the limitations of observing methods. Prior studies on  $\text{CH}_4$  flux in coastal vegetated ecosystems have mostly used the chamber-based methods, which normally cover daytime-only measurements at the risk of missing large  $\text{CH}_4$  emissions at night (Podgrajsek et al. 2014; Y. P. Zhang, Qin, et al. 2022; X. Zhu, Chen, et al. 2024). Meanwhile, the  $\text{CH}_4$  budget is typically assessed using short-term field measurements on a few consecutive days, which is prone to bias the aggregation when summed up to monthly or even longer time scales (Hou et al. 2016; Y. P. Zhang, Qin, et al. 2022; Roth et al. 2023). In addition,  $\text{CH}_4$  emission via the pulse ebullition might not be covered by the short-term chamber measurements, which could lead to an obvious underestimation of  $\text{CH}_4$  emission (Al-Haj and Fulweiler 2020). Furthermore, different anthropogenic managements across aquaculture stages can also significantly affect the magnitude and temporal variability of

$\text{CH}_4$  flux (Y. P. Zhang, Qin, et al. 2022; Xiong et al. 2024). With the rapid turnover of macroalgae biomass and sediment accumulation rates (Attard et al. 2019; Hurd et al. 2022), the macroalgae aquaculture areas may behave as biogeochemical hotspots of air-sea  $\text{CH}_4$  exchange (Lundevall-Zara et al. 2021; Hall et al. 2025).

The  $\text{CH}_4$  flux derived from the high-frequency and continuous eddy covariance (EC) measurements represents the net  $\text{CH}_4$  flux as a result of the processes of  $\text{CH}_4$  production, consumption, and transport, compensating for the limitations of traditional chamber-based methods. As an automatic and non-destructive in situ flux measuring technique, the EC approach can quantify multiple GHG fluxes simultaneously (e.g.,  $\text{CH}_4$  and  $\text{CO}_2$ ) (X. Zhu et al. 2019; Baldocchi 2020). Although the EC approach has been widely used in various ecosystems, including coastal wetlands (J. Liu et al. 2020; Y. Zhang et al. 2023; X. Zhu, Chen, et al. 2024), it has rarely been applied to coastal macroalgae aquaculture ecosystems that are highly productive and globally distributed. In this study, we used the EC approach to measure net ecosystem exchange of  $\text{CH}_4$  and  $\text{CO}_2$  for one full year over the sea-air interface in a semi-enclosed macroalgae aquaculture bay in southeast China. The key aims of this study are (1) to examine the temporal variability of  $\text{CH}_4$  fluxes across diurnal, daily, and seasonal scales and their environmental controls, and (2) to assess the importance of  $\text{CH}_4$  fluxes in affecting the radiative forcing induced by  $\text{CO}_2$  fluxes in coastal macroalgae aquaculture ecosystems. This study will provide an important piece of direct evidence on high-frequency air-sea  $\text{CH}_4$  fluxes in macroalgae aquaculture ecosystems, which helps to improve the understanding of their GHG flux dynamics and radiative forcing across time scales.

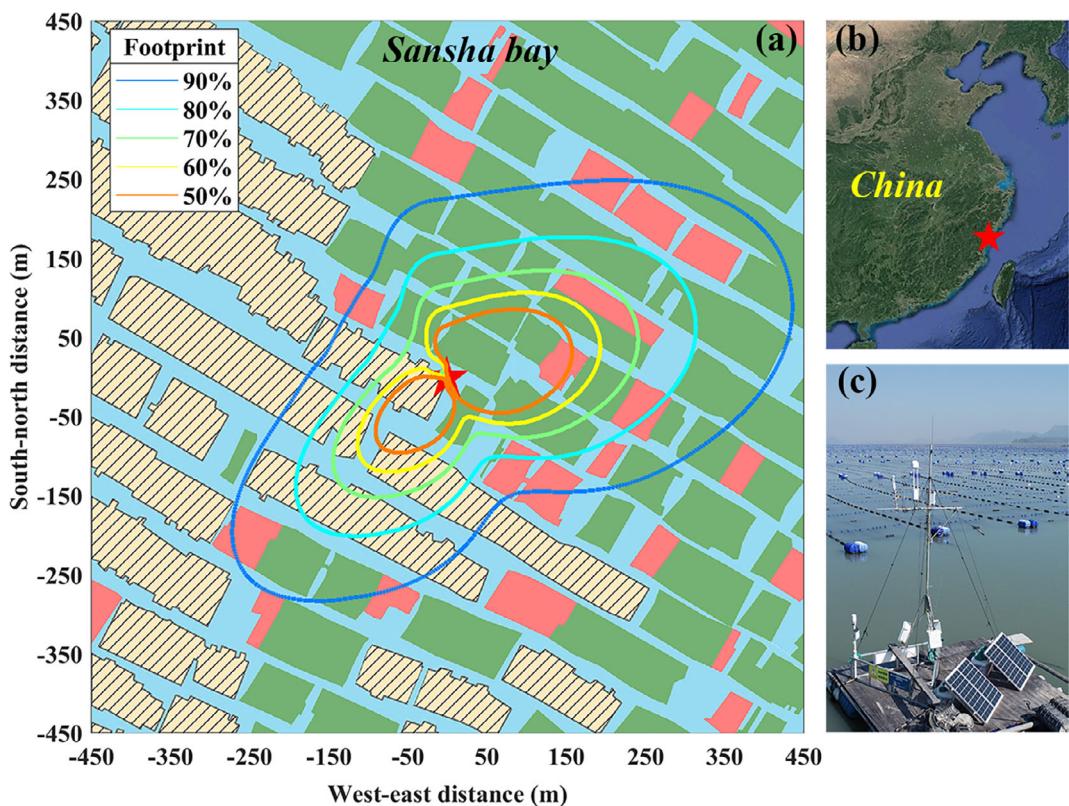
## Materials and methods

### Study area and aquaculture management

The GHG flux tower (26.7218°N, 119.9871°E; ChinaFLUX) in the study area is located in the Sansha Bay macroalgae aquaculture area in Fujian Province, China (Fig. 1), which has a subtropical monsoon climate with irregular semi-diurnal tides (mean tidal range of 5.35 m). Sansha Bay is a semi-enclosed bay with its baymouth just 3 km wide connecting to the East China Sea. *Saccharina japonica* and *Gracilaria lemaneiformis* are the two primary farming types in this macroalgae aquaculture area, with the former cultivated from December to May of the next year and the latter be cultivated all year round. There is also a portion of the fish raft around the flux tower for *Apostichopus japonicus* (sea cucumber) cultivation, with the cultivation cycle from roughly November to March of the next year.

### Greenhouse gas fluxes and ancillary measurements

The GHG fluxes (i.e.,  $\text{CH}_4$  and  $\text{CO}_2$ ) in this study referred to air-sea net ecosystem exchange of  $\text{CH}_4$  (NME) and  $\text{CO}_2$  (NEE) between the aquaculture ecosystem and the atmosphere, where



**Fig. 1.** The eddy covariance flux tower (red star) was deployed over the macroalgae aquaculture area (mapped from a drone image acquired on January 26, 2024) overlaid by the flux footprint climatology (a) in a subtropical enclosed bay in southeast China (Sansha Bay) (b). The tower was fixed to a large fish raft to reduce potential swaying issues (c). Green, red, yellow, and blue blocks indicate *Saccharina japonica*, *Gracilaria lemaneiformis*, *Apostichopus japonicus* cultivations, and seawater, respectively. Note that there was a slight spatial relocation (< 500 m apart) of the tower in mid-July 2023 due to the unintended removal of the original fish raft attached with the tower.

NME denoted the net flux of  $\text{CH}_4$  production and consumption and NEE represented the balance between  $\text{CO}_2$  influx (e.g., via gross primary production) and  $\text{CO}_2$  efflux (e.g., via ecosystem respiration). The meteorological sign convention was employed to indicate the GHG fluxes, where positive and negative values denoted upward (emission) and downward (uptake) fluxes, respectively. The NME and NEE were derived from the measurements by an EC system, consisting of two open-path gas analyzers of  $\text{CH}_4$  (Li-7700, Li-COR Inc.) and  $\text{CO}_2$  (Li-7500, Li-COR Inc.) and a three-axis sonic anemometer (CSAT-3, Campbell Scientific, Inc.) (Fig. 1c). The EC system was installed  $\sim 5$  m above the water surface of the macroalgae aquaculture area, with 90% of the fluxes contributing from the aquaculture area within  $\sim 400$  m around the tower based on the footprint climatology analyses following Kormann and Meixner (2001) (Fig. 1a). The gas analyzers' sensor mirrors were routinely cleaned approximately once a week to ensure the quality of the measurements.

An integrated weather station (ATMOS 41, METER Group, Inc.) was used to collect the meteorological data including solar radiation, wind speed, air temperature, and rainfall in this study. Water quality parameters from a nearby buoy device included water temperature, dissolved oxygen, and salinity, measured by a CTD sensor (SBE 37 SMP-ODO,

Sea-bird). Tidal level data were derived from the tidal forecast table of a nearby site within Sansha Bay (Zhangwan site from State Oceanic Administration of China). Drone and satellite imagery were combined to map the evolution of primary aquaculture types over the studying period (Fig. 1a). A high-performance Attitude and Heading Reference System Ellipse-A (SBG system) attitude system was utilized to continuously record the attitude data for 1 month (December 2023), which were used for flux corrections due to the swaying issue of the platform following Edson et al. (1998). Given that we had only 2-month attitude data and the correction only caused a 3% relative deviation, further data analyses were based on the fluxes without the attitude correction to ensure consistency.

#### Flux processing and statistical analyses

The raw 10-Hz EC data over the 1-yr studying period, from June 2023 to May 2024, were processed into 30-min time-series data using the EddyPro software (Li-COR Inc.) with a series of flux corrections (including axis rotation, frequency response correction, ultrasonic correction, and Webb-Pearman-Leuning correction) and quality control (including steady-state test, turbulent conditions test, statistical test, and absolute limits test) processes (X. Zhu, Hou, et al. 2021;

X. Zhu, Qin, et al. 2021; Y. Zhang et al. 2023; X. Wang and Zhu 2024). The 30-min data were assigned with quality flags using the 0–1–2 system (Mauder et al. 2013) (i.e., the flag of 0, 1, and 2 stands for the best-, good-, and poor-quality data, respectively), and the data with a flag of 2 were excluded in this study. The 30-min flux data under conditions of rainfall and insufficient nighttime turbulence were also discarded. Specifically, a friction velocity threshold was calculated for each 3-month time window over the year and then applied to exclude nighttime fluxes below the friction velocity threshold (Papale et al. 2006). To prevent low-quality data from possible contamination of Li-7700 mirrors, NME data with a relative signal strength indicator of less than 20% were also eliminated. Due to the unbalanced data availability between daytime and nighttime, we, respectively, performed quality control for daytime and nighttime data to mitigate the bias in the aggregation to daily and monthly data (Y. Liu and Zhu 2024). The days with valid 30-min records less than 25% of the total number of records were not incorporated into the daily and monthly scales (X. Zhu, Sun, et al. 2021; Z. Zhu and Zhu 2025). Annual NME and NEE were estimated by summarizing their respective monthly values. Over the 1-yr study period, after excluding poor data due to quality issues and instrument failures, the percentages of valid 30-min NME and NEE data were 65.7% and 66.2%, respectively. The percentages of valid daily NME and NEE accounted for 83.1% and 92.9%, respectively. To match the EC data, all ancillary measurements were converted into 30-min time series data for further analyses.

To better analyze NME across various farming periods and stages, we separated the 1-yr studying period into the following periods based on the dominant farming activities (Table 1): from December 2023 to May 2024 for *S. japonica* farming period (December and January for the rapid growth stage, February and March for the middle growth stage, and April and May for the harvesting stage); June, July, October, and November 2023 for *G. lemaneiformis* farming period. August and September 2023 were considered as non-farming (seawater) periods without macroalgae cultivation. Mean diurnal variation analyses were used to examine diurnal variation in NME. To examine the environmental controls on NME, principal component analysis and Spearman's correlation analysis were applied to 30-min time series data of NME and environmental factors. To compare the radiative forcing of NME ( $RF_{CH_4}$ ) and NEE ( $RF_{CO_2}$ ) and their net combination ( $RF_{net}$ ), we converted them into  $CO_2$  equivalent ( $CO_2\text{-eq.}$ ) using the SGWP metric

(Neubauer and Megonigal 2015), where the conversion value of 45 was used for  $CH_4$  for a 100-yr time horizon (Eq. 1).

$$RF_{net} = RF_{CH_4} \times 45 + RF_{CO_2} \quad (1)$$

Positive and negative radiative forcing values represented warming and cooling effects, respectively, from NME, NEE, or their net combination. To explore the tidal impacts, we followed our previous study (Deng et al. 2025) to divide the 30-min data even into six groups by tidal level (i.e., 1H–6H denoted six equal parts from low to high tidal levels) and treated 1H and 6H (3H and 4H) as the “slow” (“fast”) tidal current group. All data processing and statistical analyses were performed in MATLAB R2024b software (MathWorks, Inc.).

## Results

### Temporal variations in environmental factors

The 1-yr measurements showed that daily and monthly environmental factors in the study area exhibited strong seasonal variations (Fig. 2). Daily cumulative solar radiation ranged from 3.1 to  $76.6 \text{ MJ m}^{-2} \text{ d}^{-1}$ , with maximum solar radiation occurring in July. Throughout the year, daily wind speed had strong fluctuation with an average value of  $3.1 \text{ m s}^{-1}$ . Rainfall was concentrated in August and September, with the maximum daily rainfall of 134 mm over the year. Daily air temperature ranged between  $4.2^\circ\text{C}$  and  $33.2^\circ\text{C}$ , while daily water temperature ranged between  $12.8^\circ\text{C}$  and  $30.3^\circ\text{C}$ . Water quality factors also showed large seasonal fluctuations, with mean dissolved oxygen and salinity of  $6.4 \text{ mg L}^{-1}$  ( $4.1\text{--}8.4 \text{ mg L}^{-1}$ ) and 29.8 ppt (28.3–32.4 ppt), respectively. Daily maximum tidal level ranged from 6.1 to 8.8 m with peaking monthly spring tide in October.

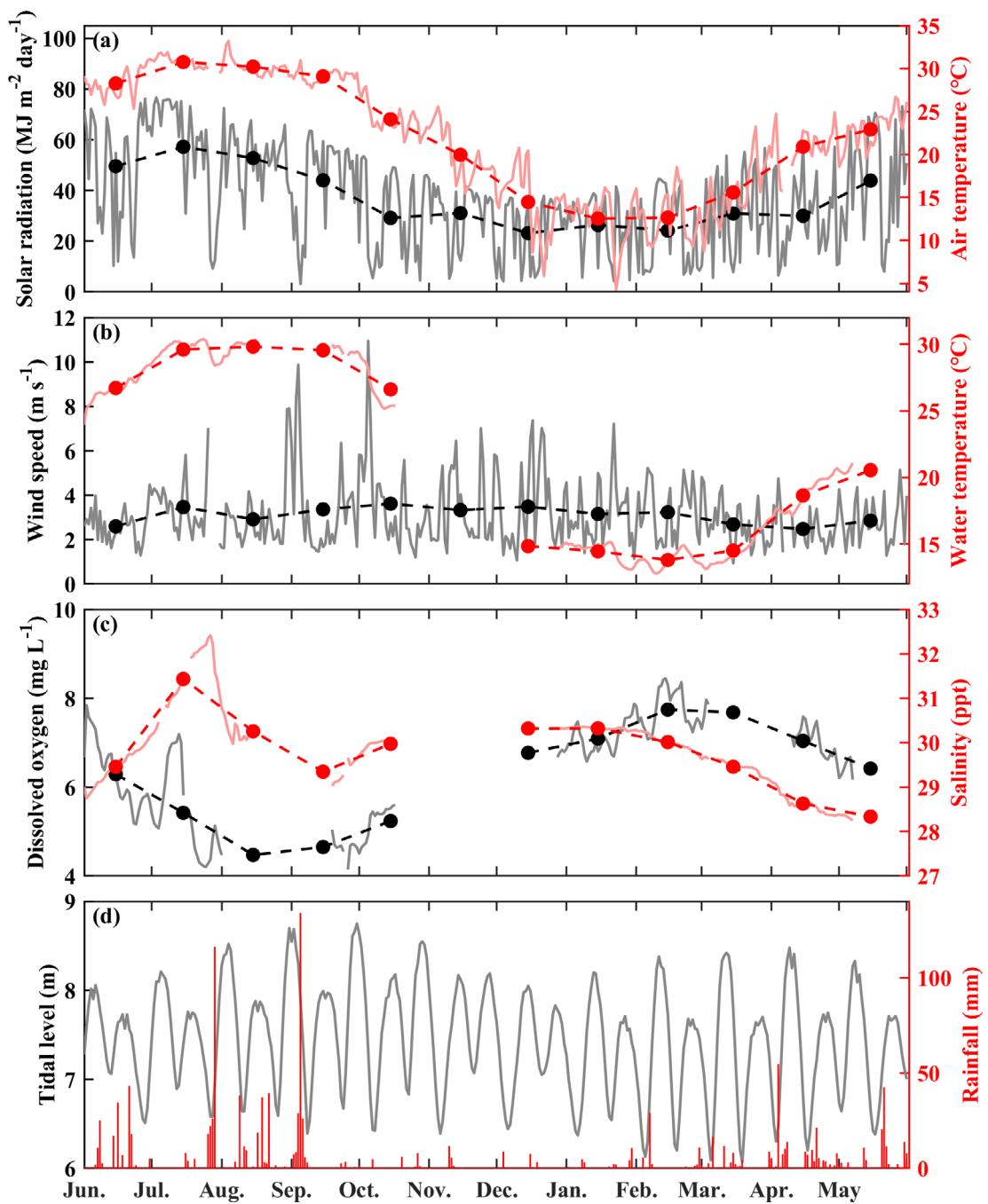
### Temporal variations in GHG fluxes

Large fluctuations in GHG fluxes were also observed on daily and monthly scales (Fig. 3). Daily NEE ranged from  $-2.3$  to  $1.5 \text{ g C m}^{-2} \text{ d}^{-1}$  with the mean value of  $0.2 \text{ g C m}^{-2} \text{ d}^{-1}$ , and monthly NEE varied from the largest source of  $18.7 \text{ g C m}^{-2} \text{ month}^{-1}$  in August to the largest sink of  $-13.5 \text{ g C m}^{-2} \text{ month}^{-1}$  in December. Over the year, this ecosystem acted as a  $CO_2$  source of  $59.8 \text{ g C m}^{-2} \text{ yr}^{-1}$ . Daily  $CH_4$  flux ranged from  $-7.3$  to  $9.2 \text{ mg C m}^{-2} \text{ d}^{-1}$  (mean value of  $1.1 \text{ mg C m}^{-2} \text{ d}^{-1}$ ). For most of the study period, the macroalgae aquaculture area acted as a net source of  $CH_4$  with the largest emission of  $142.1 \text{ mg C m}^{-2} \text{ month}^{-1}$  in July. Overall,

**Table 1.** Different periods with the dominant farming activities from June 2023 to May 2024.

2023						2024					
Jun	Jul	Aug	Sep	Oct	Nov	Dec	Jan	Feb	Mar	Apr	May
<i>Gracilaria</i> <i>lemaneiformis</i>		Non-farming		<i>Gracilaria</i> <i>lemaneiformis</i>		<i>Saccharina</i> <i>japonica</i>					
						Rapid growth		Middle growth			Harvesting

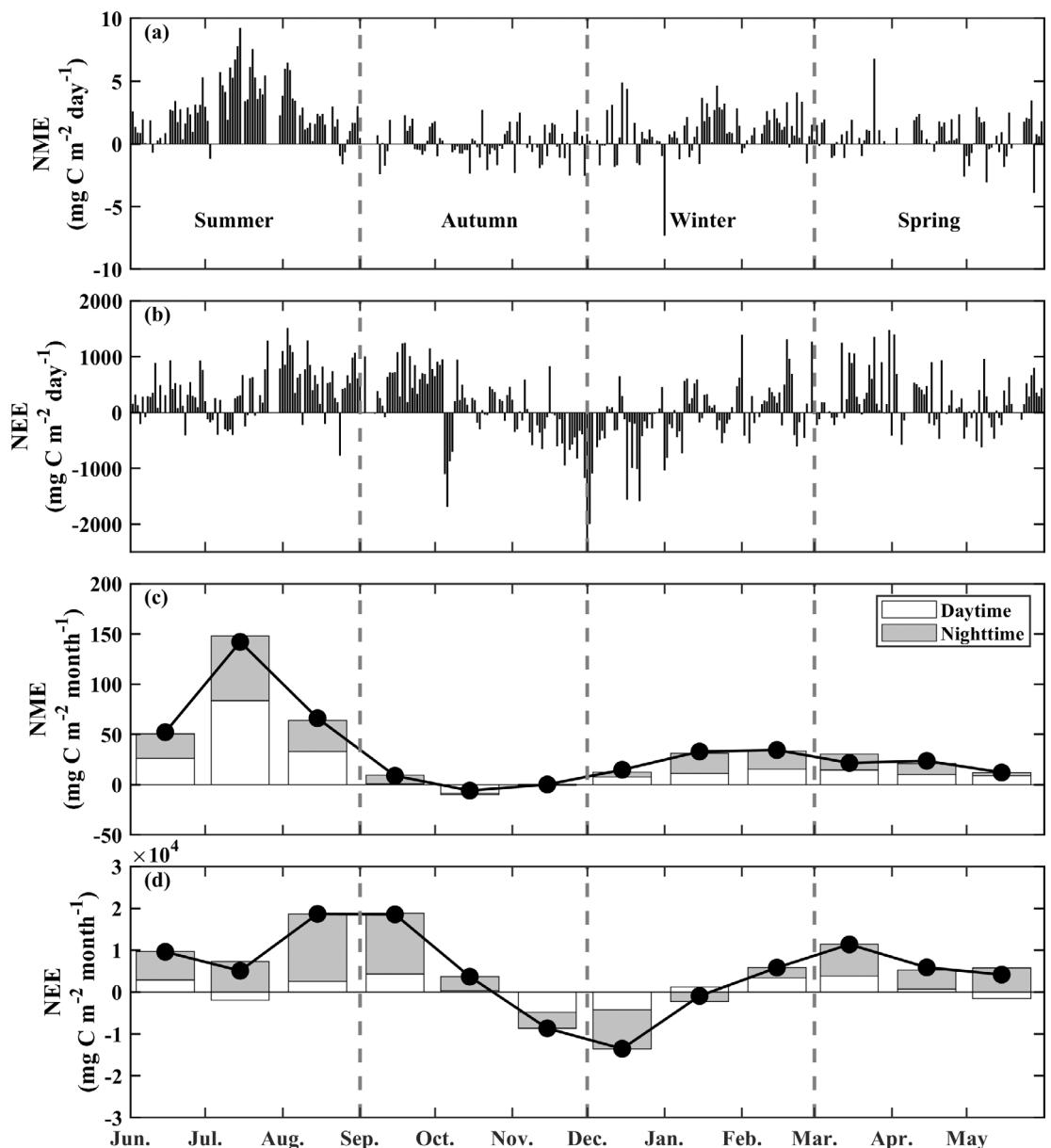
Note: “Non-farming” and Dec 2023 ~ May 2024 for “*Saccharina japonica*”.



**Fig. 2.** Temporal variations in daily (lines/bar) and monthly (solid dots) environmental factors from June 2023 to May 2024, including (a) cumulative solar radiation, mean air temperature, (b) mean wind speed and mean water temperature, (c) mean dissolved oxygen, mean salinity, (d) maximum tidal level and cumulative rainfall.

summer had the strongest  $\text{CH}_4$  emission with an average daily flux of  $2.69 \text{ mg C m}^{-2} \text{ d}^{-1}$ , whereas autumn had the lowest  $\text{CH}_4$  emission at  $0.0057 \text{ mg C m}^{-2} \text{ d}^{-1}$  (see Supporting Information Fig. S1 for the comparisons of daytime, nighttime, and all-day NME among seasons). The mean diurnal variation analyses also confirmed the  $\text{CH}_4$  emission source-sink pattern in this macroalgae aquaculture area (Fig. 4). The magnitude of

the fluctuation in mean diurnal fluxes varied across seasons with the largest fluctuation occurring in spring from  $-1.99$  to  $3.16 \text{ mmol m}^{-2} \text{ s}^{-1}$ . All of the highest NMEs averaged over the daytime ( $2.45 \text{ mmol m}^{-2} \text{ s}^{-1}$ ), nighttime ( $2.79 \text{ mmol m}^{-2} \text{ s}^{-1}$ ), and all-day ( $2.59 \text{ mmol m}^{-2} \text{ s}^{-1}$ ) periods occurred in summer. There was no statistically significant difference in NME between daytime and nighttime periods (see Supporting Information



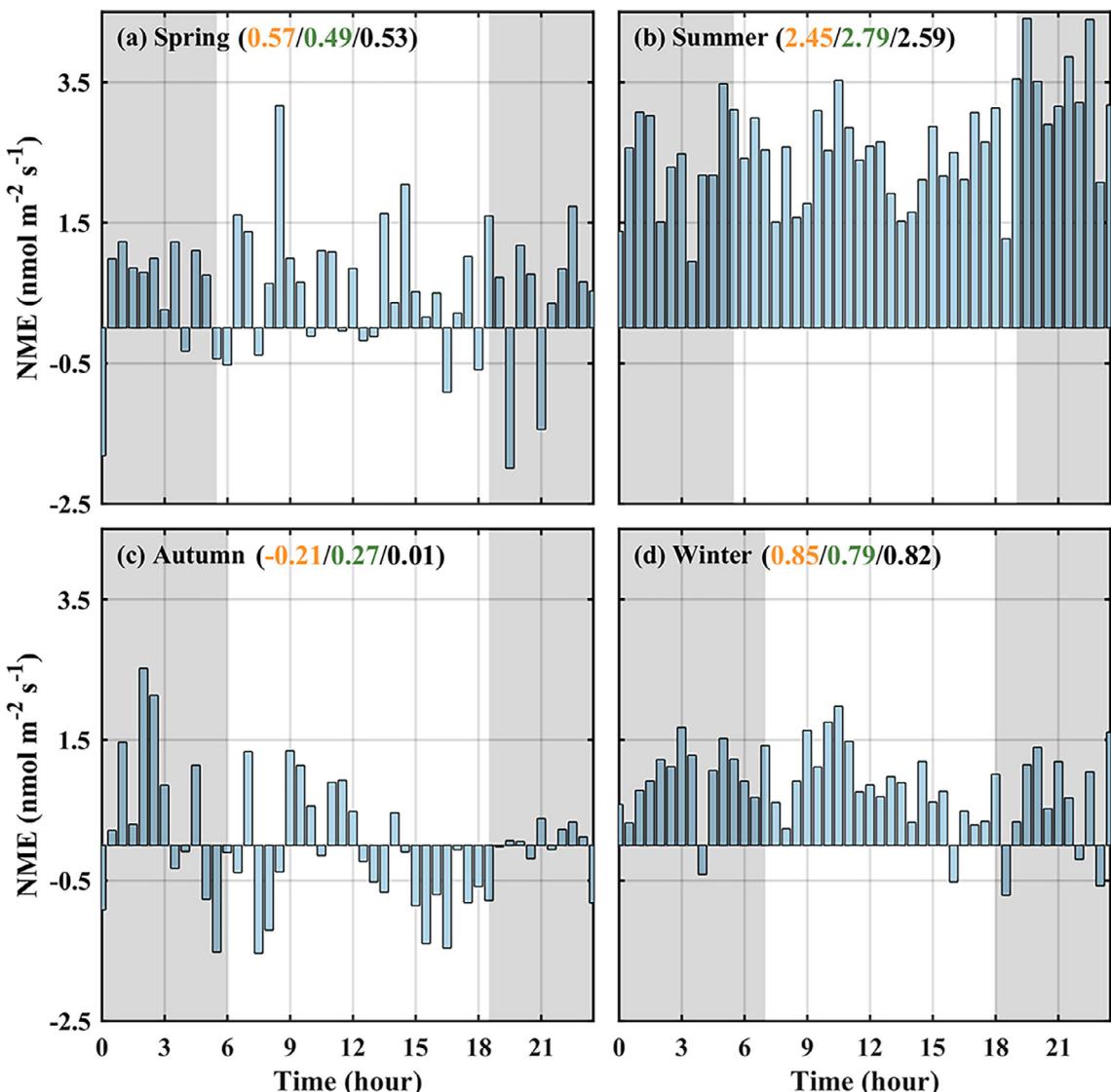
**Fig. 3.** Temporal variations in (a, b) daily and (c, d) monthly net ecosystem exchanges of  $\text{CH}_4$  (NME) and  $\text{CO}_2$  (NEE) over the macroalgae aquaculture from June 2023 to May 2024. NEE, net ecosystem exchange; NME, net ecosystem methane exchange.

Fig. S2 for monthly comparisons of daytime and nighttime NME).

#### Environmental controls on $\text{CH}_4$ fluxes

To better examine the relationships between NME and environmental factors, principal component analysis was used to explore the correlation between 30-min NME and environmental factors across different seasons and farming periods (Fig. 5). The analysis indicated that the initial two principal components explained 55.9% of the variance within the dataset. Specially, the primary component accounted for 36.3% of the variance with significant contribution from air

temperature (54.3%), water temperature (56.6%), and dissolved oxygen (−52.2%). Observations aligned with the first principal component showed pronounced separation across winter, spring, and summer/autumn, indicating the influence of temperature on seasonal  $\text{CH}_4$  flux dynamics. NME also showed a significant segregation between the kelp aquaculture period and other stages. The second principal component captured 19.6% of the variance, primarily driven by water level (64.7%). Spearman's rank correlation analysis was also applied to examine the relationships between 30-min NME and air temperature, water temperature, dissolved oxygen and water level across seasons (Fig. 6). The analysis indicated that the



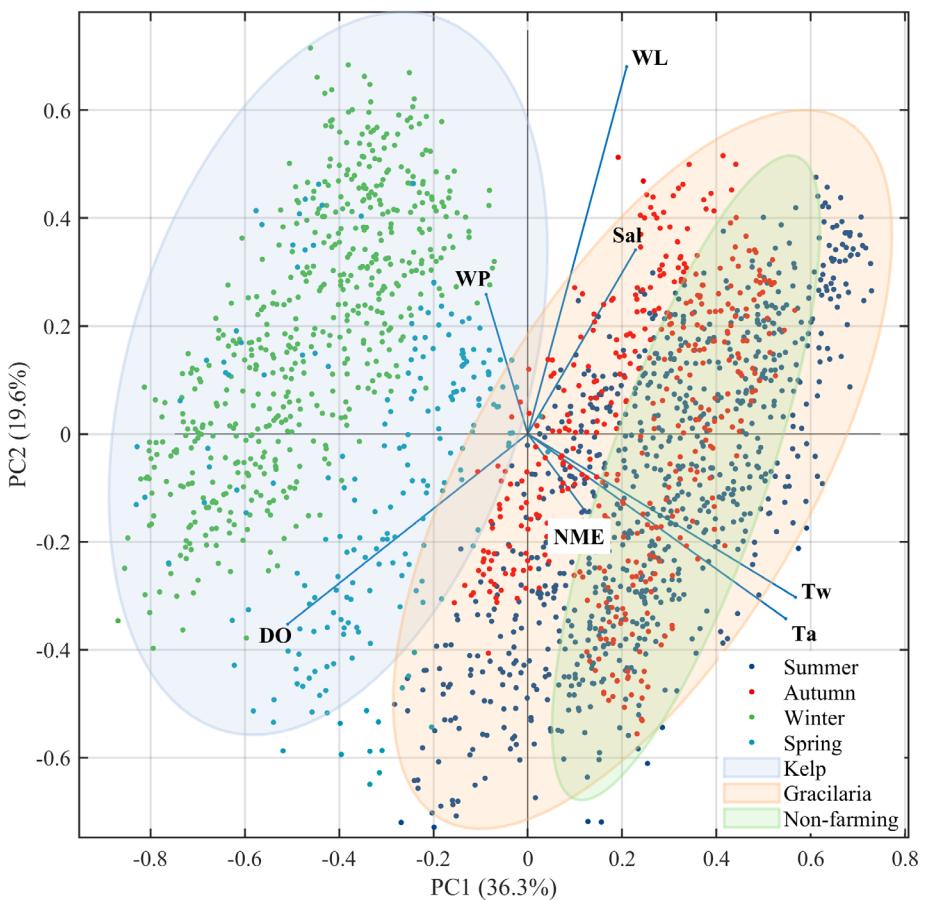
**Fig. 4.** Mean diurnal variations in net ecosystem exchange of  $\text{CH}_4$  (NME) over the macroalgae aquaculture for each season. Daytime and nighttime are indicated by white and gray areas, respectively. The three values shown in the bracket for each season indicate the mean fluxes during the daytime, nighttime, and throughout the day based on the mean diurnal variation. NME, net ecosystem methane exchange.

best explanatory factor of NME changed across seasons. Net ecosystem methane exchange was best correlated with air/water temperature during spring and summer (statistically significant positive correlations at  $p < 0.05$ ), while during autumn and winter, water level served as the best explanatory factor (statistically significant positive correlations). Dissolved oxygen exhibited a statistically significant negative correlation with NME only in summer. To further explore the links of GHG fluxes and tidal activities, the data were grouped according to different tidal levels (Fig. 7). The  $\text{RF}_{\text{CH}_4}$ ,  $\text{RF}_{\text{CO}_2}$ , and their combination of  $\text{RF}_{\text{net}}$  tended to show bimodal varying patterns during both daytime and nighttime, having stronger radiative forcing values at faster tidal currents (when tidal level changed most rapidly; i.e., around 3H–5H) (see

Supporting Information Fig. S3 for statistical comparisons among all tidal levels). The percentage of  $\text{RF}_{\text{CH}_4}$  over  $\text{RF}_{\text{CO}_2}$  ranged from 1.7% to 25.4% (excluding an extreme value with tiny NEE) across tidal levels with overall higher daytime values.

#### Temporal variations in global warming potentials

The relative strength of NME and NEE expressed as SGWP varied across seasons and farming stages for this macroalgae aquaculture area (Fig. 8). The  $\text{CH}_4$  emission increased the warming effect caused by  $\text{CO}_2$  emission throughout all seasons except winter. The largest and smallest enhancements occurred in summer (12.2%) and autumn (0.8%), respectively, while the  $\text{CH}_4$  emission counteracted 44.2% of the cooling



**Fig. 5.** Principal component analysis of the relationships between net ecosystem exchange of  $\text{CH}_4$  (NME) and environmental factors, including air temperature ( $T_a$ ), water temperature ( $T_w$ ), water level (WL), dissolved oxygen (DO), and wind speed (WP). This biplot showed the distribution and the loadings of all factors along the first two principal components (PC1 and PC2). Each vector represents a factor with the direction and magnitude reflecting the contribution of the factor to each principal component axis. Close (symmetrical) vectors indicate a positive (negative) correlation, while orthogonal vectors indicate no correlation between the factors. NME, net ecosystem methane exchange.

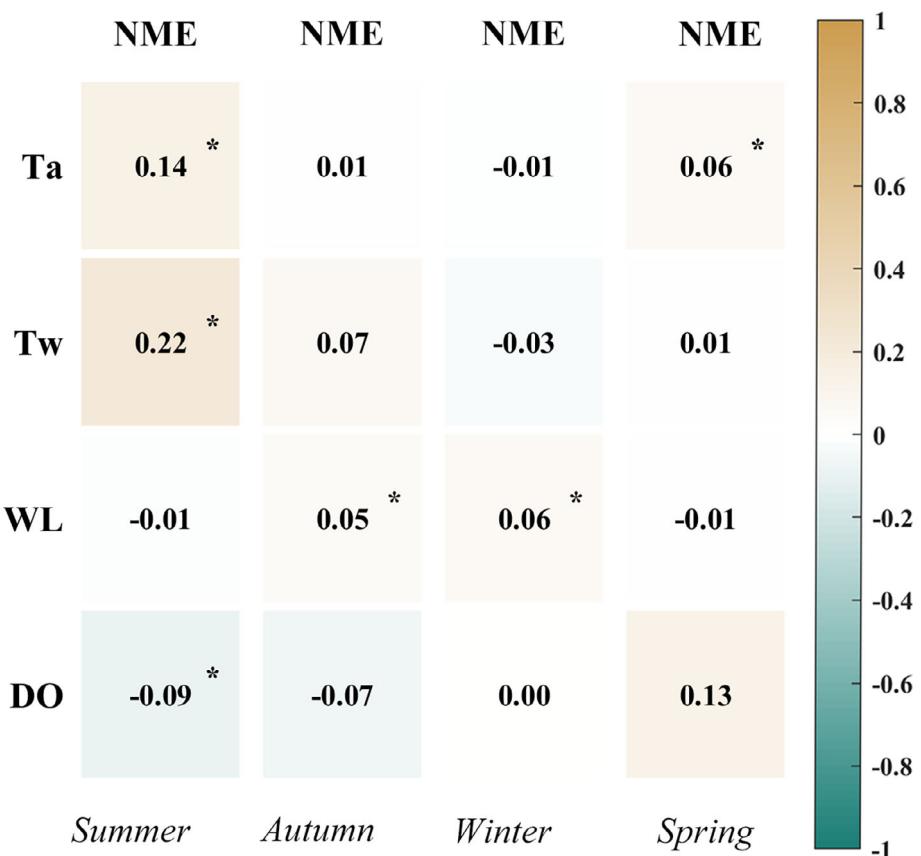
effect due to  $\text{CO}_2$  uptake in winter. In terms of farming periods, NME intensified the NEE-induced warming effect with comparable enhancements between kelp and *Gracilaria* farming periods and smaller enhancement for the non-farming period. During the rapid growth stage of kelp farming, NME offset 9.8% of the NEE-induced cooling effect, while NME enhanced the NEE-induced warming effect with comparable percentages between the middle growth and harvesting stages. On an annual basis, the  $\text{CH}_4$  emission intensified the NEE-induced warming effect with a 11.0% increase, promoting the net radiative forcing from 219.3 to 243.4  $\text{g CO}_2\text{-eq m}^{-2} \text{yr}^{-1}$ .

## Discussion

### $\text{CH}_4$ emission and its temporal variations

One-year high-frequency and continuous flux observations show that the macroalgae aquaculture area acts as a  $\text{CH}_4$  source in most months with an annual emission of

0.40  $\text{g C yr}^{-1}$ , which confirms the role of macroalgae aquaculture areas as potential  $\text{CH}_4$  emission hotspots. This annual emission is comparable to  $\text{CH}_4$  emission from natural macroalgae habitats in the Baltic Sea (0.25  $\text{g C yr}^{-1}$ ; Roth et al. 2023) and a macroalgae aquaculture area in Sanggou Bay (0.21  $\text{g C yr}^{-1}$ ; Hou et al. 2016), but lower than  $\text{CH}_4$  emission from the coastal mariculture ponds (1.44  $\text{g C yr}^{-1}$ ; Y. F. Zhang, Tang, et al. 2022). In comparison with traditional blue carbon ecosystems, this annual emission is lower than the median annual emission in mangrove (1.22  $\text{g C yr}^{-1}$ ) and saltmarsh (0.98  $\text{g C yr}^{-1}$ ) but slightly higher than that in seagrass (0.28  $\text{g C yr}^{-1}$ ) (Al-Haj and Fulweiler 2020). Similar to many other coastal ecosystems (McNicol et al. 2017; Roth et al. 2023; Henriksson et al. 2024), large seasonal fluctuations occur in  $\text{CH}_4$  fluxes in the study area. The peaking  $\text{CH}_4$  emission in summer is presumably linked to enhanced methanogenesis at higher temperatures over the year (Yvon-Durocher et al. 2014). In addition to temperature, the  $\text{CH}_4$  emission could also be



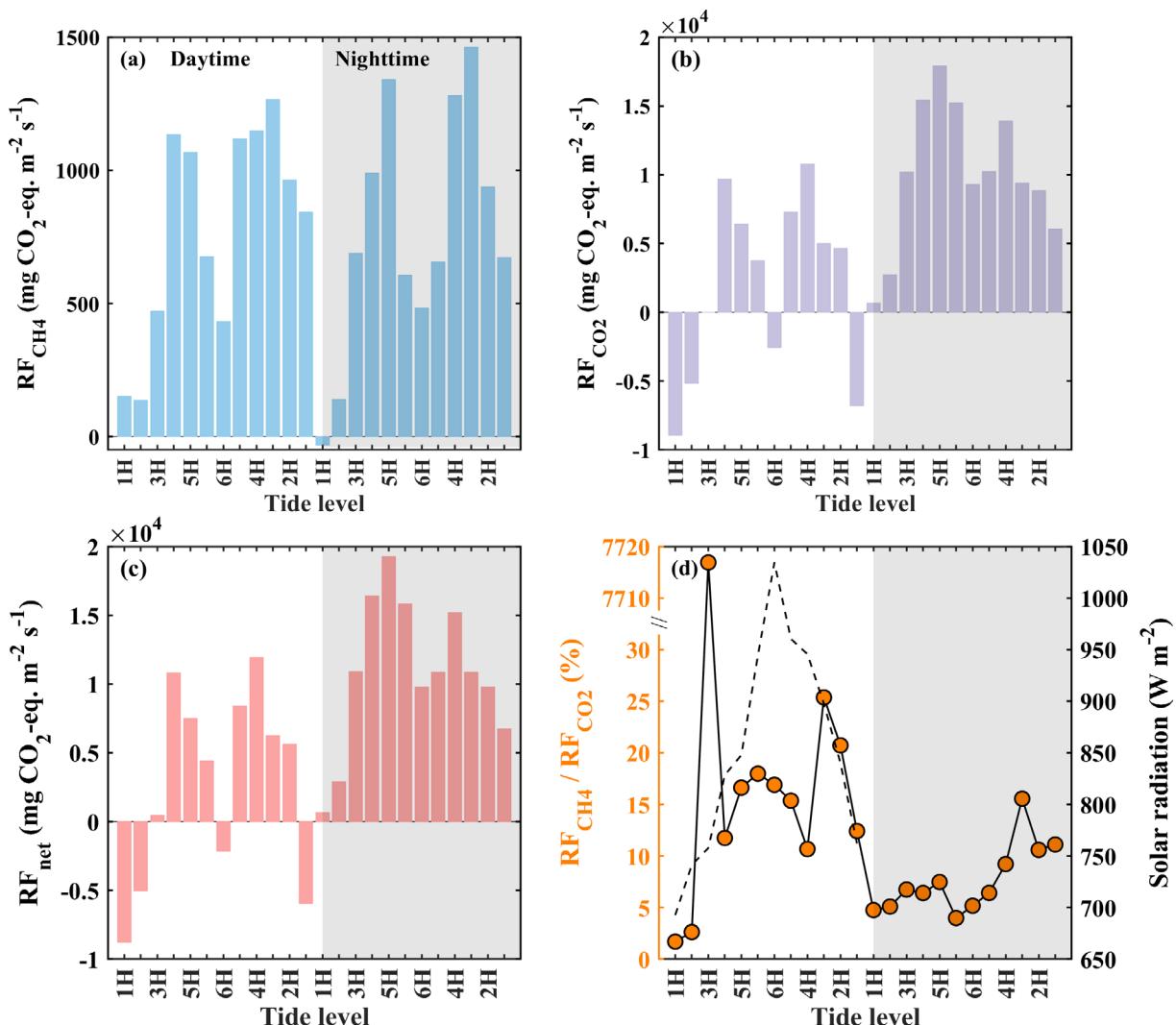
**Fig. 6.** Heatmaps of Spearman's rank correlations between 30-min net ecosystem exchange of  $\text{CH}_4$  (NME) and environmental factors, including air temperature (Ta), water temperature ( $T_w$ ), water level (WL), and dissolved oxygen (DO) across different seasons. The correlation coefficients were calculated on a pairwise basis, and the values with a star sign indicated statistical significance at  $p < 0.05$ . NME, net ecosystem methane exchange.

obviously affected by the substrate availability (Wallenius et al. 2021). For example, the second strongest  $\text{CH}_4$  emission in this macroalgae aquaculture area is observed in winter instead of autumn that has the second highest temperatures (Figs. 2, 3). This may be due to the fact that the kelp farming in winter provides more substrate inputs for methanogenesis (Y. Wang et al. 2023; Pessarrodona et al. 2024), which in turn accelerates the production and release of  $\text{CH}_4$  despite the lower temperatures (Dale et al. 2019). In terms of sub-daily variation,  $\text{CH}_4$  flux exhibits no consistent diel variation with only marginal differences observed between daytime and nighttime emission. This small day–night difference could be explained by counteractive processes. On the one hand, daytime emission can be either promoted by stronger methanogenesis with higher temperatures and elevated labile substrates derived from photosynthesis (Bridgham et al. 2013). On the other hand,  $\text{CH}_4$  emission can also be suppressed by stronger methanotrophy with more oxygen availability (or oxygen priming) from photosynthesis (Noyce et al. 2023). These counteractive effects are partially reflected by the correlation analysis showing that  $\text{CH}_4$  emission is positively correlated with temperature in summer and spring but

negatively correlated with dissolved oxygen in summer (Fig. 6).

#### Tidal controls in regulating $\text{CH}_4$ fluxes

Coastal vegetated habitats are experiencing fast environmental changes and thus the GHG flux dynamics are susceptible to a variety of factors such as the physical forcing drivers (e.g., tides), especially in the open coastal environment (Al-Haj and Fulweiler 2020; Roth et al. 2022; Arias-Ortiz et al. 2024; X. Zhu, Chen, et al. 2024). Bimodal flux variation patterns with tidal levels are observed in this study area, i.e., larger fluxes correspond to faster tidal currents when tidal level changes most rapidly (e.g., 3H–5H) (Fig. 7). This can be explained by the fact that faster tidal currents facilitate water column mixing and air–sea gas exchange. On the one hand, stronger water column mixing with sediment resuspension triggered by faster tidal currents may substantially enhance  $\text{CH}_4$  production and bubble ebullition (Jordan et al. 2020), because microorganism concentrations in the sediments can exceed those in the water column by several orders of magnitude (Schmale et al. 2015). On the other hand, faster tidal currents with well-mixing turbulent conditions in surface water



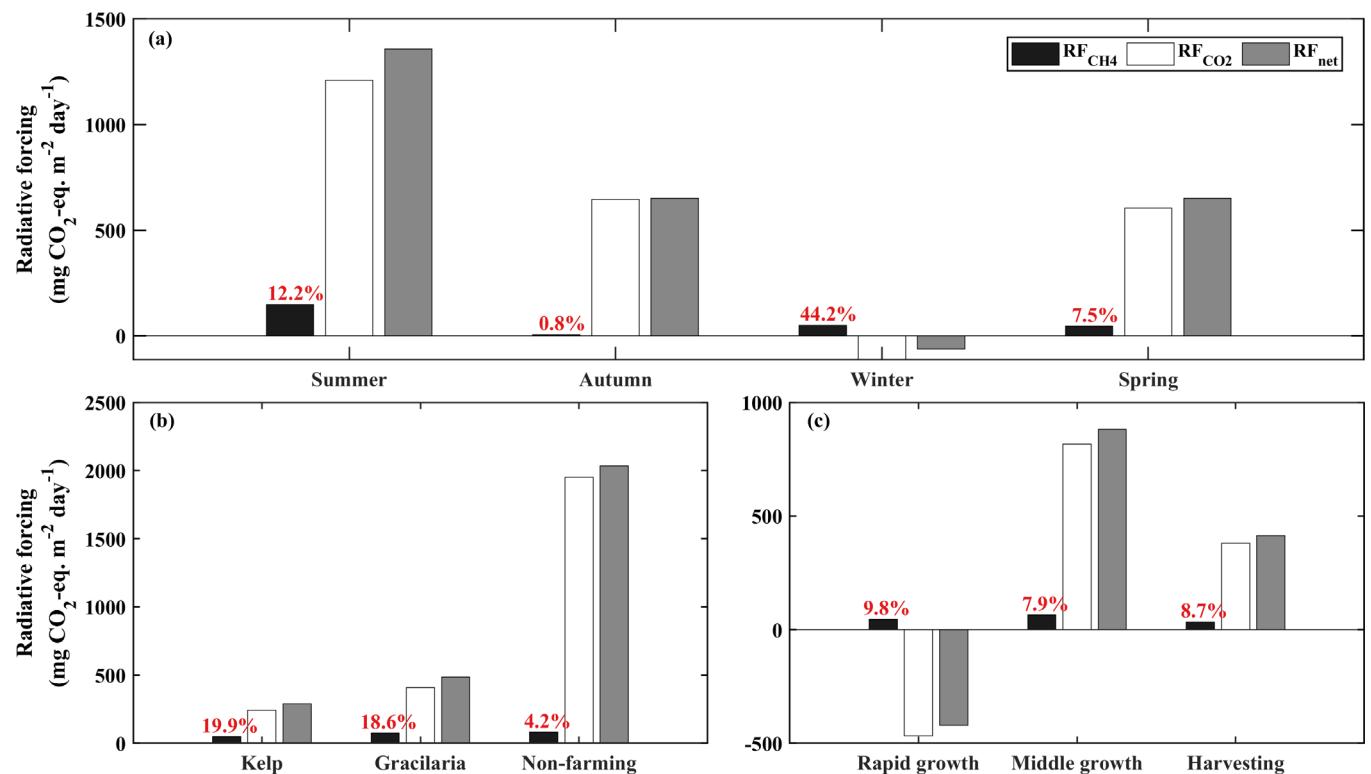
**Fig. 7.** Mean tidal variations in radiative forcing, using the sustained global warming potential metric, of (a) net CH<sub>4</sub> flux ( $RF_{CH_4}$ ), (b) net CO<sub>2</sub> flux ( $RF_{CO_2}$ ), (c) the net combination of  $RF_{CH_4}$  and  $RF_{CO_2}$  ( $RF_{net}$ ), as well as (d) the percentage of  $RF_{CH_4}$  over  $RF_{CO_2}$  (overlaid by solar radiation) across tidal levels from low to high tide (1H–6H) for both daytime and nighttime periods. The tidal rule of 12<sup>th</sup>s was used to divide tidal levels into 12 groups from 1H (lower levels) to 6H (higher levels) and from 6H to 1H within a flood-ebb tidal cycle.

can substantially promote gas transfer velocity and thus speed up CH<sub>4</sub> emission into the atmosphere, since gas transfer velocity is a key control of gas exchange across the air-sea interface (Wanninkhof et al. 2009). Given that peaking solar radiation at noon roughly matches the highest tidal levels at daytime (Fig. 7d), the low value of the bimodal pattern at noon could also be associated with stronger methanotrophy with more oxygen supply from midday photosynthesis.

#### Asynchronous variations between GHG fluxes

The CH<sub>4</sub> emission exacerbates the positive radiative forcing from CO<sub>2</sub> efflux in this macroalgae aquaculture area, in contrast to the blue carbon offset effect reported by many other studies (Rosentreter et al. 2018; Roth et al. 2023; Yau et al. 2023; Li et al. 2024; X. Zhu, Chen, et al. 2024). The net

radiative forcing varies considerably across different seasons and farming phases, with the strongest percentage contribution of CH<sub>4</sub> over CO<sub>2</sub> in winter rather than in summer with peaking CH<sub>4</sub> emission. This is mainly due to the asynchronous variations between CH<sub>4</sub> and CO<sub>2</sub> fluxes that are also reported in other coastal ecosystems (Roth et al. 2023; X. Zhu, Chen, et al. 2024). In this study, this asynchrony might mainly result from various farming activities, for example, kelp farming in winter not only promotes photosynthetic carbon uptake leading to a weaker CO<sub>2</sub> source but also provides more substrates for methanogenesis. Furthermore, when farming activities are not conducted, CO<sub>2</sub> emission increases dramatically, leading to a much smaller enhancement in the warming effect from CH<sub>4</sub> emission. The fluctuations in the percentage contribution of CH<sub>4</sub> over CO<sub>2</sub> can vary by one or two orders of



**Fig. 8.** Mean radiative forcing of net ecosystem exchanges of CH<sub>4</sub> (RF<sub>CH4</sub>) and CO<sub>2</sub> (RF<sub>CO2</sub>) as well as their net combination (RF<sub>net</sub>) across various (a) seasons, (b) farming periods, and (c) growth stages of kelp farming. The radiative forcing values were quantified as CO<sub>2</sub> equivalent (CO<sub>2</sub>-eq.) using the sustained-flux global warming potential metric for a 100-yr time horizon. The value above each RF<sub>CH4</sub> bar indicates the percentage of RF<sub>CH4</sub> over RF<sub>CO2</sub>.

magnitude across seasons, farming stages, and tidal levels (Figs. 7, 8). Despite being relatively low in the contribution of CH<sub>4</sub> emission on an annual scale, it should be taken into account in the assessments of radiative forcing in macroalgae aquaculture ecosystems. This is especially important during critical periods such as CO<sub>2</sub> depletion or low emission periods (Yang et al. 2018), when CH<sub>4</sub> emission may account for a large fraction of the total GHG budget.

### Limitations and uncertainties

The field measurements and related data analyses of this study suffer from several limitations and uncertainties. First, the flux data unavoidably become biased as a result of flux corrections and quality control. For example, to reduce the uncertainty due to unbalanced day-night data availability, we have to first, respectively, perform quality control on daytime and nighttime data and then aggregate them into the daily and monthly data (X. Zhu, Ma, et al. 2024). Second, coastal ecosystems are susceptible to the terrestrial water inflows, which could favor a stronger source of CH<sub>4</sub> (Borges et al. 2018). However, given there is no river running into the aquaculture area where the EC tower is located, the confounding effects of background water inflows might be relatively small. Third, due to the high cost of the construction of the EC system in such a bay habitat experiencing severe weather (e.g., typhoons), we

cannot afford to conduct paired EC measurements. Here we only analyzed and compared the impact of farming activities from the perspective of farming cycle without a control setup. Future research could use other approaches such as the floating chamber to improve the flux analysis regarding spatial heterogeneity. Finally, although the SGWP metric is used to evaluate the warming effect of CH<sub>4</sub> emission in this study, this metric only represents a specific effect over a fixed time period (Ma et al. 2024). Future application of dynamic SGWP metrics should improve the assessment of the radiative forcing of GHG fluxes at different time scales (Lynch et al. 2020).

### Conclusions

One-year EC measurements of CH<sub>4</sub> fluxes over a subtropical macroalgae aquaculture ecosystem in southeast China were used in this study to examine the temporal variability of CH<sub>4</sub> fluxes across time scales and its contribution to the radiative forcing of CO<sub>2</sub> expressed as SGWPs. We confirm this ecosystem acts as a net CH<sub>4</sub> source on an annual scale with the flux fluctuations mainly driven by temperatures and tidal activities among seasons. Furthermore, for a 100-yr time horizon, annual CH<sub>4</sub> emission leads to a one-tenth increase in the warming effect from annual CO<sub>2</sub> emission. This is the first EC study to disentangle the temporal variability of NME and its

radiative contribution over macroalgae aquaculture ecosystems based on simultaneous measurements of both CH<sub>4</sub> and CO<sub>2</sub> fluxes. Although the ecosystem is confirmed as a biogeochemical hotspot of both CH<sub>4</sub> and CO<sub>2</sub> fluxes, their asynchronous flux variations complicate the assessments of their net radiative forcing. Future studies with multi-year EC measurements of continuous and high-frequency GHG fluxes are highly needed to further reduce the uncertainty in the assessments of GHG budgets and their radiative forcing in macroalgae aquaculture ecosystems.

## Author Contributions

Yueteng Deng: writing – original draft, visualization, investigation, data curation, formal analysis. Xianghui Guo: writing – review and editing, resources. Dengjin Hu: investigation. Hui Luo: visualization, investigation. Yougan Chen: writing – review and editing. Xudong Zhu: writing – review and editing, resources, supervision, project administration, funding acquisition, conceptualization.

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## Conflicts of Interest

None declared.

## Data Availability Statement

Data will be made available on request.

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### Supporting Information

Additional Supporting Information may be found in the online version of this article.

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